

# NASA Technical Memorandum 100708

## Data Report for the Siple Coast Project

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OCTOBER 1988

(NASA-TM-100708) DATA REPORT FOR THE SIPLE  
COAST (ANTARCTICA) PROJECT (NASA) 108 P  
CSC1 08B

N89-10403

Unclassified  
G3/43 0168978





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National Aeronautics and  
Space Administration

Scientific and Technical  
Information Office

1988



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## INTRODUCTION

This report contains a compilation of the data collected by a combined NASA/University of Chicago field program on Siple Coast, Antarctica from October 1983 to January 1987 and a description of the methods used to obtain the data. In its first three field seasons, beginning in October 1983, the project has concentrated on studying Ice Streams B and C (see Figure 1). A number of publications have resulted from the analysis of these data and they are referenced in Appendix 4. This report will assist in the continuing analysis of these data and adds them to the growing data base that exists for the region.

The NASA/University of Chicago program is part of a larger project which also involves the Byrd Polar Research Center at Ohio State University and the Geophysical and Polar Research Center at the University of Wisconsin. The major goals of this larger project are:

- 1) To measure the mass balance and configuration of the major ice streams in the Ross Sea drainage basin of the West Antarctic ice sheet;
- 2) To determine the forces that control the flow of these ice streams;
- 3) To investigate long-term trends in the drainage basin.

To meet these goals, the NASA/University of Chicago group designed a data acquisition operation with the following specific objectives:

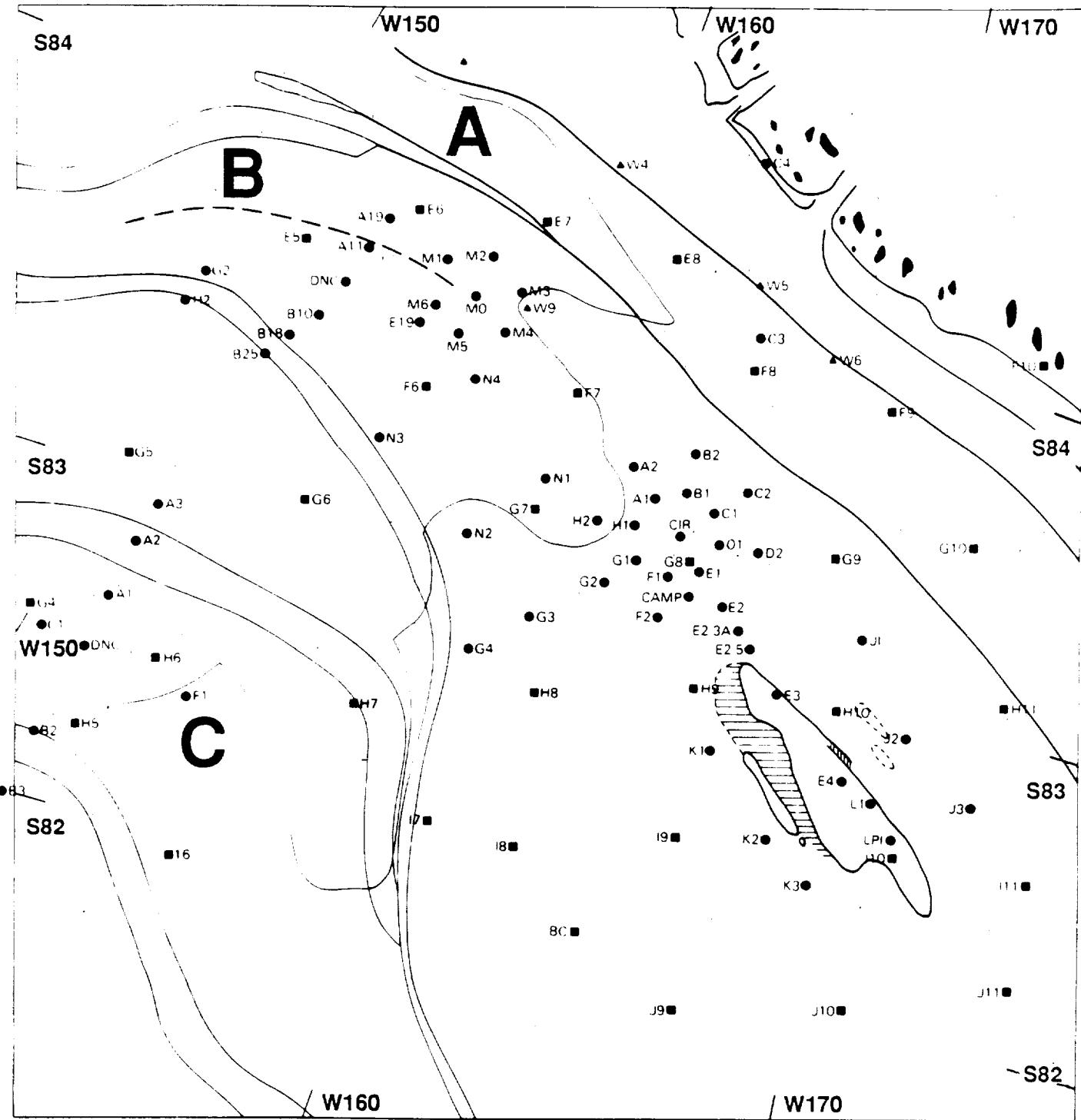
- 1) Establish a network of stations to measure the surface velocity and strain rate profiles across the widths of Ice Streams B and C at locations near their grounding lines to determine net ice stream discharge rates.

- 2) Establish stake schemes to measure longitudinal velocity and strain rate profiles to determine the spatial gradients of both ice transport and the forces which restrain the ice stream motion.
- 3) Establish a network of stations which surround Crary Ice Rise (Figure 1), to measure velocities and strain rates, and thus determine its influence on the flow of Ice Stream B, and the state of mass balance of the ice rise.
- 4) Establish a network of stations which surround a feature named Ice Rise "a", a feature seen on early airborne photography in the region, to determine its position and characteristics. (We state here that Ice Rise "a" is not a true ice rise at all, as its velocity is the same as the surrounding ice, which is itself, lightly grounded.)
- 5) Establish a network of stations to measure the ice velocity and strain rate fields in the lower, lightly grounded regions of Ice Stream B referred to as the "ice plain".
- 6) Map the precise location of the grounding lines of Ice Streams B and C.
- 7) Perform regional studies of small-scale ice ripples to characterize their effect on the large-scale flow.
- 8) Establish stake schemes and carry out resection surveys to measure the velocity field in the severely crevassed margins of both ice streams and ice rises, to determine the local stresses in this region.
- 9) Develop a new method for rapid acquisition of strain rate data using multi-leg rosettes.

Although the data presented within this report are mainly the result of the collaborative effort of the University of Chicago

and NASA, the wider collaboration with Ohio State University and the University of Wisconsin is also evident, particularly in the map figures. We have used the results published in Shabtaie and Bentley (1986) as the basis for the boundaries of the ice streams and their grounding lines and the radar definition of ice rises and ripples. We are also grateful for vigorous continuing discussions between the authors and the other collaborating institutes; in particular, with S. Shabtaie, C. R. Bentley and I. M. Whillans.

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**Figure 1: Regional map showing station positions.**

## SECTION 1: ICE VELOCITY

During this program, 26 ice velocities were measured by accurately determining the position of stakes using a Doppler satellite tracking system at several times and then calculating the distance moved during the time intervals. To obtain the distance between two geographic positions, Clarke's formulae were used, (see Jackson 1980), with a value for the semi-major axis of 6378.16 km and a flattening of 1/298.25.

A Magnavox MX1502 geoceiver was used for precise position determination. Each successful tracking of a satellite pass provides an independent fix of the receiver position. Successive fixes are averaged together by the receiver so that after about 30 passes, the position reaches a steady value within a sphere of about a 1.6-meter radius. However, this value of the standard deviation cannot be considered the error in absolute position because there is also an error in the broadcast satellite ephemeris which can cause an additional error of up to 20 meters (Whillans, personal communication). This imprecision can be improved with post-processing by using either precise ephemerides or translocation techniques. To obtain 30 passes at 84 degrees south, the average latitude of the field program, the geoceiver must operate for about 24 hours. For operational reasons, a geoceiver was rarely allowed to track more than 40 passes. More often, only about 10 passes were used to obtain a position.

We have adopted the error analysis of Thomas et al (1984) in which for less than 30 passes, they take the root-mean-square radial error as  $17.48/\sqrt{n}$  meters, where n is the number of passes used to obtain a point position. This is twice the error found during the geoceiver test program and about 1.5 times larger than the error calculated by the Magnavox MX1502, but we feel this larger error should be used, given the lack of any other control. This error indicates that for ice moving at a velocity greater than 150 m/a, and a measurement interval of at least 1 year, as

few as 10 passes will give the velocity to within 5 percent; 30 passes will reduce the error to less than 3 percent. For ice moving 100 m/a, 30 point positions are needed to give 5 percent accuracy, while for ice moving at 30 m/a, translocation methods are needed to obtain better than 10 percent accuracy.

Table 1 summarizes the velocities measured by geoceivers. It includes the latitude and longitude of each station, the date at which the original position was observed, the number of passes for the two point position determinations, and the velocity magnitude and direction along with an estimate of their errors. Our computed velocities and those of the Ross Ice Shelf Geophysical and Glaciological Survey (RIGGS) program are plotted in Figure 2.

Table 2 gives the velocity across a 54-km width of Ice Stream B, including station DNB. The width of the glacier at this point is  $84 \pm 3$  km taken from the map in Shabtaie and Bentley (1987). The velocities were determined by transecting a line of stakes twice to determine their relative positions and motion, and then adding the velocity determined by Doppler satellite positioning at station DNB. The observations were made over two 1-year time intervals; stakes A11 to B10 were surveyed in 1983 and 1984, and stakes B11 to B19 were surveyed in 1984 and 1985 so all the observations were reduced to two epochs. The rotation of the stake line (about 6 minutes in 2 years) was included in these calculations. The velocity at stake B18 was also determined by Doppler satellite positioning. The misclosure between the two methods is 5 meters in magnitude and 5 degrees in azimuth, which corresponds to a 26-meter error for both positions each time they were measured. Figure 3 plots the transverse velocity profile.

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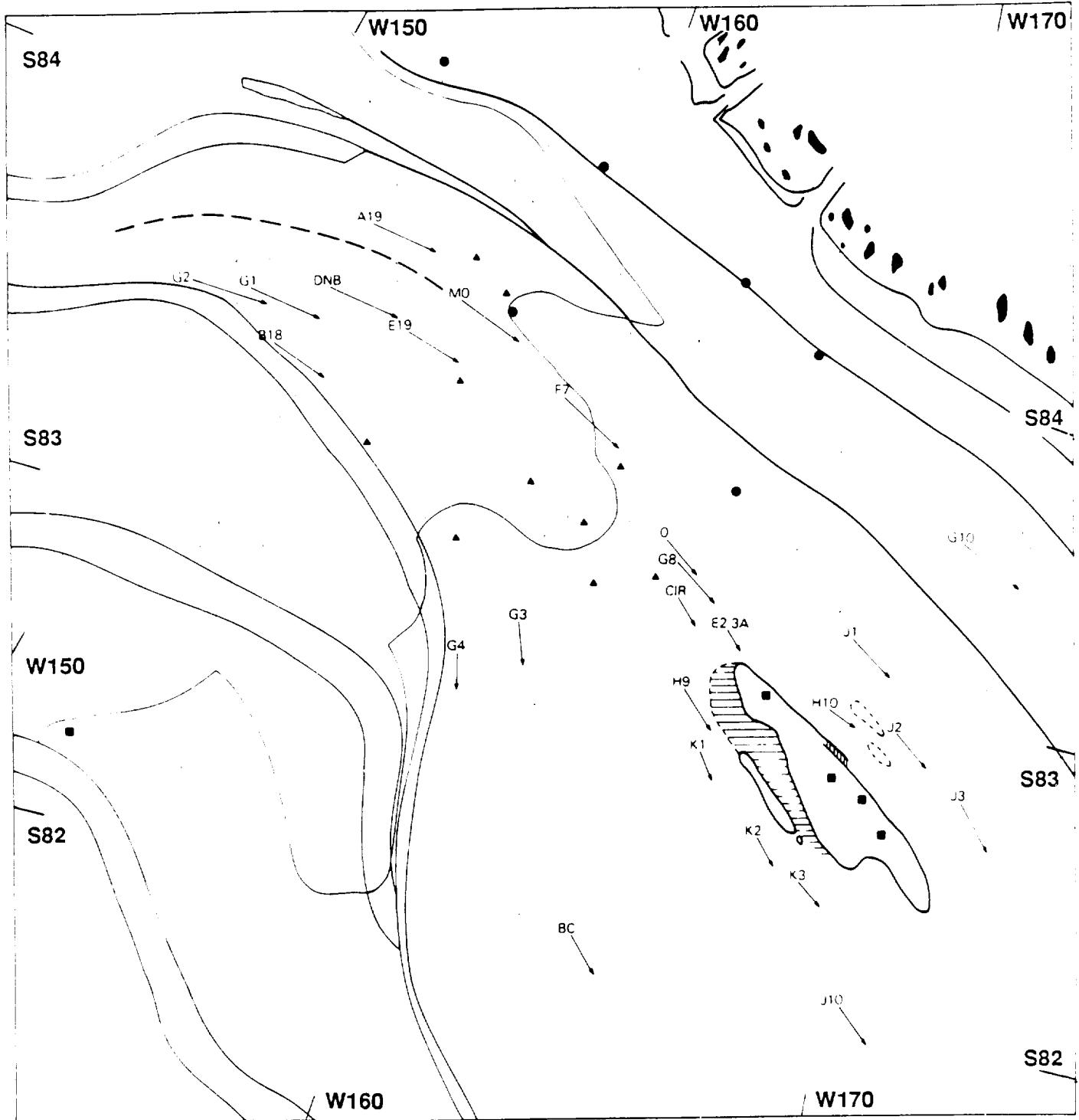


Figure 2: Regional map showing station velocities.

TABLE 1 VELOCITY MEASUREMENTS

Site	Latitude	Longitude	No. of Passes	Velocity			Date
				YR1	YR2	(m/a)	
	deg/min/sec	deg/min/sec				(deg)	D/M/Y
<b>Crary</b>							
E2.3A	83 33 09.1	168 12 43	17 9	172 + 7		319 + 2	30.1 11 84
E3	83 22 44.0	169 34 33	6 30	21 8		144 - 21	19.7 11 84
E4	83 9 20.2	171 36 38	15 7	7 8		23 49	19.9 11 84
K1	83 10 24.3	168 9 9	5 8	225 10		326 3	20.1 11 84
K2	82 56 34.9	169 58 00	7 8	224 9		323 2	20.2 11 84
K3	82 49 26.6	171 09 52	13 4	239 10		313 2	25 11 84
J1	83 35 26.7	171 35 34	6 11	387 9		310 1	20.4 11 84
J2	83 19 3.8	173 04 26	3 15	333 11		314 2	20 11 84
J3	83 7 22.4	174 55 29	18 16	385 6		329 1	27.9 11 84
C4	84 57 52.1	165 38 44	1 14	75 18		25 13	23.0 11 84
G3	83 25 42.3	162 43 12	5 9	312 10		338 2	21 11 84
G4	83 16 40.0	161 26 55	2 20	264 14		342 3	23.2 11 84
L1	83 06 11.8	172 25 21	25 40	15 5		333 17	26.2 11 84
LP1	83 00 04.3	172 59 57	9 20	16 7		126 24	28.1 11 84
O	83 47 14.0	166 01 28	8 10	369 8		304 1	30.9 11 84
CAMP	83 37 14.4	166 44 31	17 28	245 5		318 1	17.0 11 84
<b>Downstream B</b>							
MO	84 17 45.9	158 10 58	12 23	471 6		285 1	4 12 84
DNB	84 10 27.6	154 18 43	45 23	517 5		269.5 0.5	19.4 12 84
E19	84 9 3.1	156 51 07	5 38	465 8		279 1	7.2 12 84
G1	84 3 57.6	152 08 59	5 4	534 12		269 1	12.3 12 84
G2	84 0 28.2	150 32 03	17 11	551 7		259.9 0.7	15.9 12 84
A19	84 24 26.5	154 42 21	15 20	487 6		270.6 0.7	11.2 12 84
B18	83 57 00.1	153 33 32	12 23	426 6		279.1 0.8	9.0 12 84
B25	83 52 01.8	153 10 38	4 23	13 9		268 38	9.3 12 84
H2	83 53 49.0	150 25 25	19 38	30 5		203 10	17.3 12 84
<b>Downstream C</b>							
H5	82 35 25	153 14 54	27	6 .5		321 7	12 12 84

\*All azimuths are measured clockwise from true North

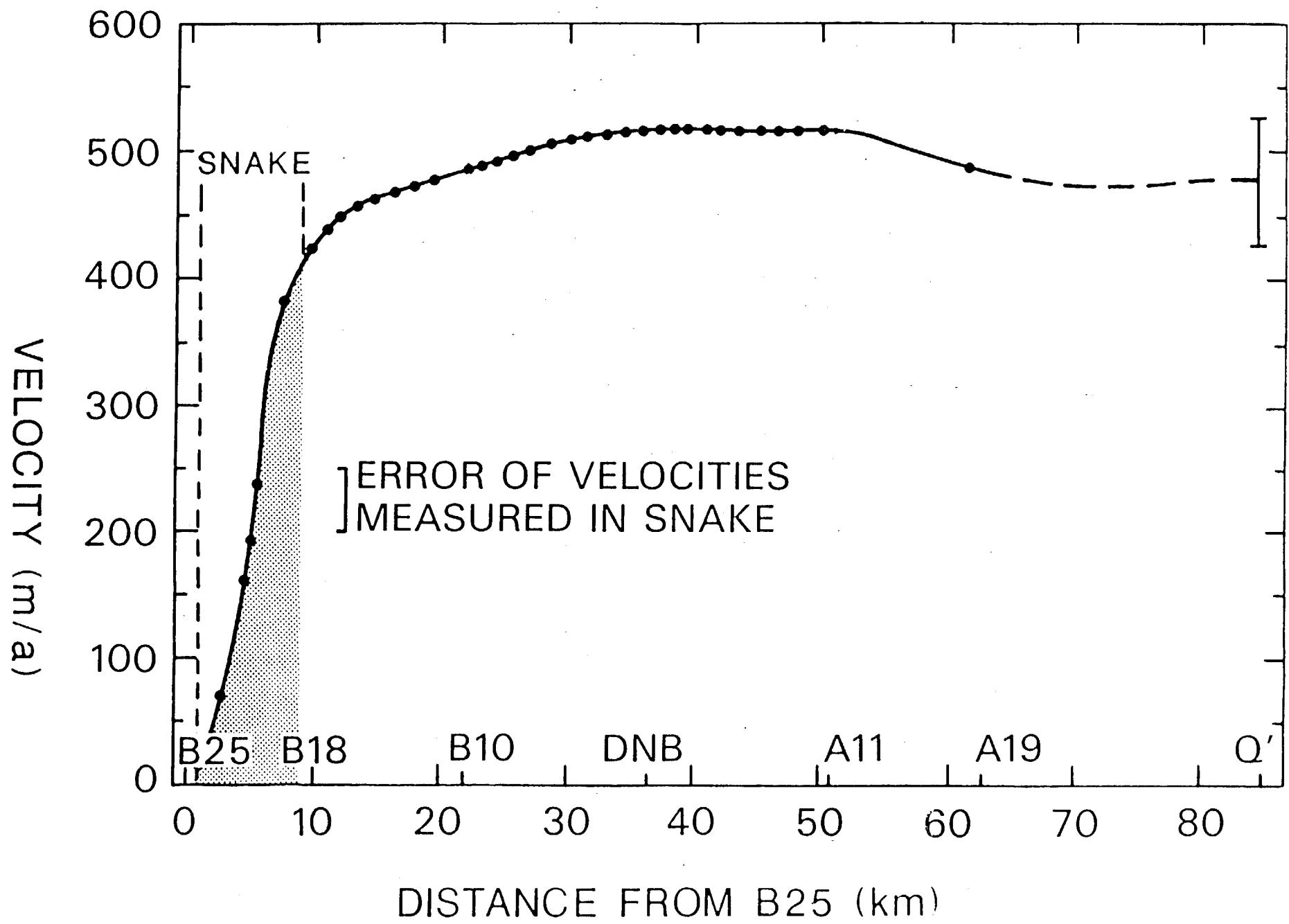
TABLE 2 VELOCITIES ACROSS ICE STREAM B MEASURED BY TRANSECT

STATN	LAT	LONG	X	Y	VELOCITY			
					deg/min/sec	deg/min/sec	km	km
B19	83 55 58.1	153 24 01	301.68	602.45	386.1	0.6	274.3	210.9
B18	83 56 56.9	153 30 11	299.79	601.36	431.2	0.5	274.3	210.8
B17	83 57 38.0	153 34 07	298.54	600.57	445.9	0.5	274.3	210.7
B16	83 58 18.6	153 38 04	297.29	599.79	455.4	0.4	274.1	210.5
B15	83 58 59.9	153 42 06	296.02	598.99	462.3	0.4	274.0	210.3
B14	83 59 42.8	153 46 20	294.69	598.17	467.6	0.4	273.9	210.1
B13	84 00 21.9	153 50 13	293.49	597.41	471.8	0.4	273.6	209.8
B12	84 01 03.9	153 54 24	292.19	596.61	475.9	0.3	273.5	209.6
B11	84 01 45.6	153 58 33	290.90	595.80	480.0	0.3	273.3	209.3
B10	84 02 42.2	154 04 15	289.15	594.71	485.6	0.3	273.1	210.0
B9	84 03 29.5	154 05 25	288.31	593.49	490.4	0.2	272.7	208.6
B8	84 04 14.9	154 06 32	287.50	592.32	494.7	0.2	272.5	208.3
B7	84 04 59.8	154 07 40	286.70	591.17	498.8	0.2	272.0	207.9
B6	84 05 46.0	154 08 49	285.88	589.98	502.2	0.2	271.6	207.5
B5	84 06 32.9	154 10 00	285.05	588.78	506.2	0.2	271.2	207.0
B4	84 07 19.8	154 11 11	284.21	587.57	509.8	0.1	270.7	206.5
B3	84 08 06.2	154 12 22	283.39	586.38	512.9	0.1	270.3	206.1
B2	84 08 52.5	154 13 33	282.56	585.19	515.0	0.1	270.0	205.8
B1	84 09 41.2	154 14 48	281.70	583.93	516.2	0.1	269.8	205.5
DNB			280.87	582.74	517.0	0.1	269.0	
A1	84 11 10.3	154 17 06	280.11	581.64	517.7	0.1	269.6	205.3
A2	84 11 56.7	154 18 19	279.28	580.45	518.0	0.1	269.4	205.1
A3	84 12 41.7	154 19 28	278.49	579.29	517.8	0.1	269.0	204.7
A4	84 13 25.6	154 20 37	277.71	578.16	518.1	0.2	268.7	204.3
A5	84 14 12.3	154 21 50	276.88	576.96	517.4	0.2	268.3	204.0
A6	84 14 45.4	154 22 43	276.29	576.10	517.1	0.2	268.1	203.7
A7	84 15 31.9	154 23 56	275.46	574.91	516.2	0.2	267.9	203.5
A8	84 16 02.2	154 24 44	274.92	574.18	515.7	0.2	267.8	203.4
A9	84 16 49.0	154 26 04	274.08	572.93	516.0	0.3	267.7	203.3
A10	84 17 19.3	154 26 57	273.53	572.15	516.8	0.3	267.6	203.2
A11	84 18 05.1	154 28 15	272.70	570.98	517.7	0.3	267.7	203.3

\*Bearing is given as angle clockwise from positive X-axis.

At station A1, positive X-axis is oriented 64.3 degrees clockwise from true North.

Figure 3: Velocities across Ice Stream B.



## SECTION 2: STRAIN RATE AND ROTATION

The rate of deformation at the surface of an ice sheet with respect to mutually perpendicular axes (x, y and z) can be expressed by the sum of two second-rank tensors, assuming there is no rotation in the vertical plane,

$$\begin{pmatrix} E_{xx} & E_{xy} \\ E_{yx} & E_{yy} \end{pmatrix} + \begin{pmatrix} 0 & w_{xy} \\ -w_{yx} & 0 \end{pmatrix}$$

where the strain rate tensor is given by

$$E_{ij} = 0.5 \left( \frac{dU_i}{dj} + \frac{dU_j}{di} \right)$$

and the vorticity tensor is

$$w_{ij} = 0.5 \left( \frac{dU_i}{dj} - \frac{dU_j}{di} \right)$$

and  $U_i$  is the velocity in the  $i$  direction. The vertical strain rate is derived from the incompressibility of ice and the sum of the two principal surface strain rates:  $E_{zz} = - (E_{xx} + E_{yy})$ .

To obtain strain rates in an area, two methods were employed. Either strain rosettes or double lines of stakes were set out and surveyed twice; the second survey typically followed the initial survey after a year. In the case of strain rosettes, the strain rate is determined by assuming a linear velocity field between any three stakes in the rosette (usually the three outlying stakes). Velocities are measured relative to the central

stake, and are given by the horizontal change in position of the outer stakes divided by the time interval between surveys.

The velocity within the rosette is then expressed in the form,

$$U_x = A_1x + A_2y + A_3$$

$$U_y = B_1x + B_2y + B_3$$

where  $x$  and  $y$  are position coordinates and  $U_x$  and  $U_y$  are the velocity components in the  $x$  and  $y$  directions. The position and velocities are known for the three outlying stakes so the constants  $A_1$  to  $B_3$  can be determined from solving a set of simultaneous equations in the form,

$$\begin{pmatrix} x(a) & y(a) & 1 \\ x(b) & y(b) & 1 \\ x(c) & y(c) & 1 \end{pmatrix} \begin{pmatrix} A_1 \\ A_2 \\ A_3 \end{pmatrix} = \begin{pmatrix} U_x(a) \\ U_x(b) \\ U_x(c) \end{pmatrix}$$

where the letters in parentheses denote the three outlying stakes. There is a similar expression for the components in the  $y$  direction. Assuming infinitesimal strain, the strain rates are then given by:

$$A_1 = dU_x/dx = E_{xx}$$

$$B_2 = dU_y/dy = E_{yy}$$

$$0.5(A_2 + B_1) = 0.5(dU_x/dy + dU_y/dx) = E_{xy}$$

where positive values denote extension and negative values denote compression.

For the double line of stakes, the first stage in calculating the strain rates is again to determine the velocity field. This is

done by solving an overdetermined set of time-dependent observation equations to derive, with error estimates, the position and velocity for each stake site, using the method described in Wager et al (1980). The position and velocity of the stakes are then used in a series of interconnecting triangles down the scheme using a different linear velocity field for each triangle.

#### Measurement Errors

A combination of field measurements of distances, angles and azimuths was made to determine strain rates. During the first field season, 1983-84, distances were measured by CA 1000 Telurometers and HP Electronic Distance Meters (EDM's), which are accurate to within 3 cm per 1500 m and 1 cm per 1500 m, respectively. Geodimeter 112 EDM's were used during the next two field seasons; accurate to  $\pm$  1 cm per 1500 m. The greater accuracy of the EDM's made their horizontal positions above the stake significant. The theodolite mounting of the EDM was optically plumbed onto the stake, so we consider a reasonable error to be 1 cm. However, where strain nets were measured using only distances the misclosure was more typically 2 cm.

The mean misclosure between rounds of the angle measurements was 4 seconds of arc. However, one of the instruments in the 1984-85 season consistently misclosed to 20 or 30 seconds of arc. In this case, measurements were repeated up to five times to obtain a consistent reading. In calculating the errors of the strain rate, a minimum error of 5 seconds of arc was used; if the misclosure was greater than 5 seconds, the size of the misclosure was taken as the value of the error.

Azimuths were measured with the least accuracy. The greatest contributions to the azimuth error were in the determinations of time and longitude. A 1-second error in time contributes to about a 15-second-of-arc error in azimuth; a 1-second error in longitude gives a 1-second-of-arc error in azimuth. Time was

taken from the satellite tracking unit, or from radio time signals. Sometimes wristwatches were used which were later calibrated using a satellite tracking unit. We have not calibrated the time given by these units, but we note that when they have been run concurrently they can differ in time by up to 1 second. Therefore, we have used an error of 1 second for time.

The position in longitude is known to better than 1 minute when a geoceiver position is used, but when single SATNAV fixes are used the longitude error could be up to 10 minutes of arc. SATNAV satellite receivers (Racal-Decca 412) are a single-channel system accurate to  $\pm 200$  m after about 5 passes, and were used in the place of the MX1502 geoceivers during the 1983-84 field season. Only a few station positions have been estimated by one SATNAV position; the remainder are determined using either three SATNAV fixes or up to 50 geoceiver fixes. We have used 45 seconds of arc as the error in azimuth. This may be high for azimuths measured under optimum conditions: where time is taken directly from a geoceiver using a stopwatch to measure lapse time, and longitude is taken from a geoceiver position. Figure 4 and Table 3 present the strain rate data.

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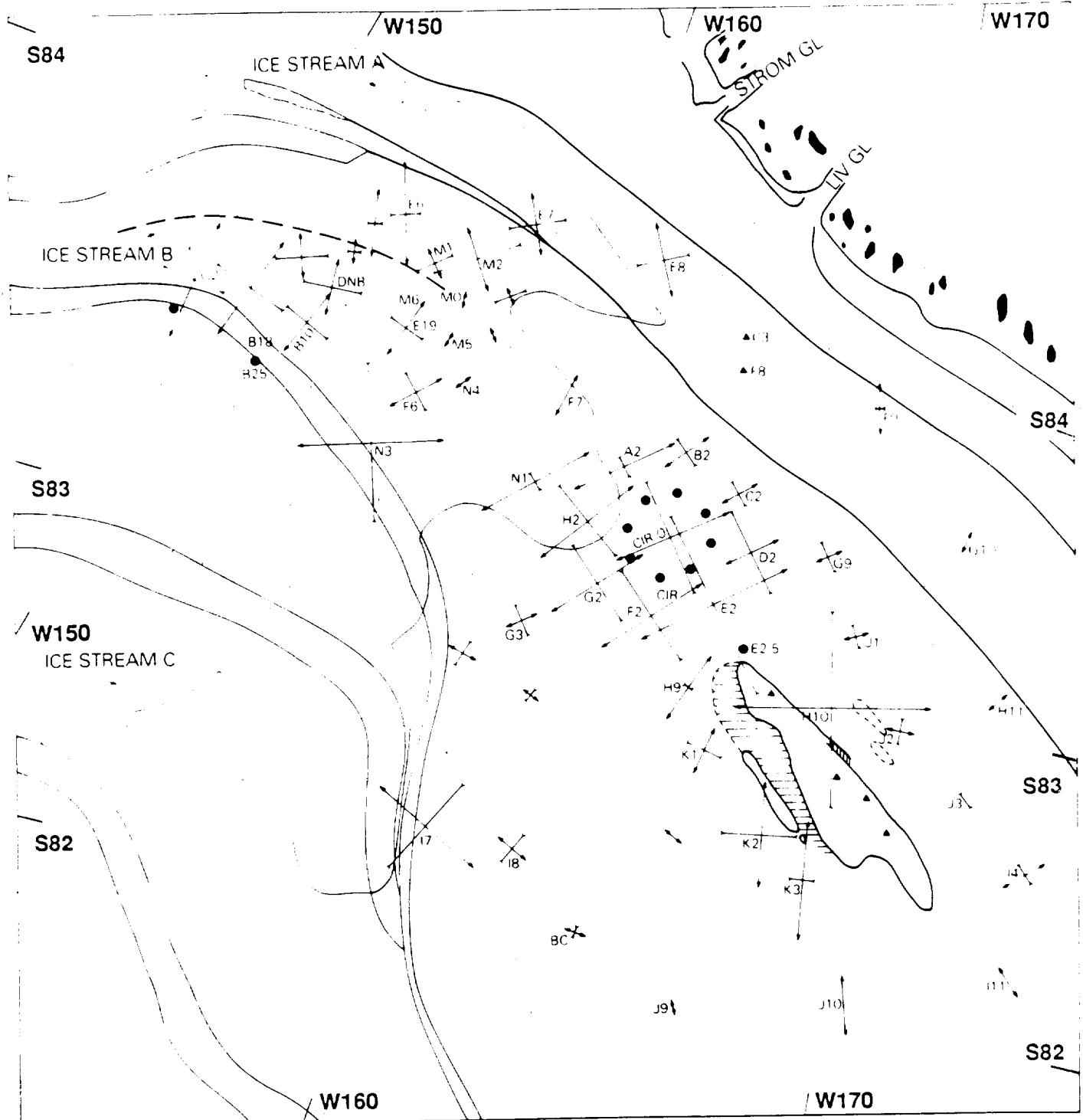


Figure 4: Regional map showing selected strain rates.

TABLE 3: STRAIN RATES

STATION	P1*	P2*	P2**	ROTATION <sup>+</sup>	BEARING <sup>†</sup>	DATE
	x10 <sup>-3</sup> a <sup>-1</sup>	x10 <sup>-3</sup> a <sup>-1</sup>	x10 <sup>-3</sup> a <sup>-1</sup>	x10 <sup>-3</sup> rad/a	of P1	
CRARY						
A1	4.04	-0.04	-0.93	-1.11	0.00	10.4
A2	3.79	0.00	0.59	2.20	0.00	15.0
B1	0.03	0.00	-0.38	-0.65	0.00	20.4
B2	1.79	0.00	0.97	0.82	0.00	25.0
C1	0.16	0.00	0.02	-0.66	0.00	30.3
C2	0.53	0.00	0.87	0.00	0.00	35.4
C3	0.04	0.00	0.02	-0.82	0.00	40.0
D1	0.53	0.00	0.02	-0.66	0.00	45.0
D2	0.97	0.00	0.02	-0.00	0.00	49.8
E1	2.49	-0.27	-0.00	0.00	0.00	54.0
E2	1.77	0.12	-0.64	1.66	0.00	59.0
E3	0.58	0.02	-0.05	0.00	0.00	64.0
E4	2.48	-0.25	-0.39	1.71	0.00	69.0
E5	1.53	0.00	-0.05	0.00	0.00	74.0
F1	0.77	0.14	-0.19	1.88	0.00	79.0
F2	1.29	0.12	-0.62	1.64	0.00	84.0
G1	0.25	0.00	-0.04	0.00	0.00	88.0
G2	0.76	0.00	-0.22	1.93	0.00	93.0
G3	0.04	0.00	-0.06	0.00	0.00	98.0
G4	0.50	0.00	-0.02	0.00	0.00	103.0
H1	0.91	0.00	-0.95	1.95	0.00	108.0
H2	1.12	0.00	-1.23	2.13	0.00	113.0
J1	0.76	0.00	-0.05	0.00	0.00	118.0
J2	0.02	0.00	-0.02	0.00	0.00	123.0
J3	0.53	0.00	-0.98	1.98	0.00	128.0
K1	0.04	0.00	-0.03	0.00	0.00	133.0
K2	0.54	0.00	-0.03	0.00	0.00	138.0
K3	0.04	0.00	-0.03	0.00	0.00	143.0
LPI	0.00	0.00	0.00	0.00	0.00	148.0
0	0.00	0.00	0.00	0.00	0.00	153.0
DOWNSTREAM B						
All	0.00	0.00	0.00	0.00	0.00	158.0
A19	2.96	0.02	-0.47	1.95	0.00	163.0
A10	2.40	0.03	-1.39	2.01	0.00	168.0
B18	0.63	0.02	-1.02	1.06	0.00	173.0
B25	2.20	0.03	-1.06	1.92	0.00	178.0
C10	1.63	0.03	-1.03	0.03	0.00	183.0
E19	3.31	0.02	-1.02	1.73	0.00	188.0
G1	2.95	0.03	-1.02	1.92	0.00	193.0
G2	1.77	0.02	-1.02	1.73	0.00	198.0
H2	1.17	0.02	-1.02	1.73	0.00	203.0
DNB	0.50	0.09	-0.02	0.04	0.00	208.0
DNB	2.05	0.09	-0.02	0.04	0.00	213.0
MO	0.00	0.48	-0.02	0.04	0.00	218.0
M1	0.32	0.01	-1.02	1.73	0.00	223.0
M2	0.32	0.02	-1.02	1.73	0.00	228.0
M3	0.32	0.04	-1.02	1.73	0.00	233.0
M4	0.52	0.03	-1.02	1.73	0.00	238.0
M5	0.52	0.02	-1.02	1.73	0.00	243.0
M6	0.52	0.06	-0.68	0.02	0.00	248.0
M7	0.27	0.06	-0.68	0.02	0.00	253.0
M8	0.11	0.07	-0.46	0.05	0.00	258.0
N1	0.56	0.05	-0.22	0.06	0.00	263.0
N2	0.07	0.06	-0.35	0.08	0.00	268.0
N3	0.11	0.07	-0.35	0.07	0.00	273.0
N4	0.56	0.05	-0.22	0.06	0.00	278.0

\* PL, P2= Principal axes of horizontal tension and compression, respectively.

\*\* PZ= -(P1+P2), vertical strain rate.

† Rotation is clockwise, and bearing is clockwise from true North.

### SECTION 3: SURFACE ELEVATIONS FROM GEOCEIVER AND OPTICAL LEVELLING

The Doppler satellite tracking system also provided measurements of surface elevation with respect to the reference ellipsoid WGS 72. This ellipsoid has a semi-major axis of 6378135 m and a flattening of 1/298.26. Table 4 gives the measured elevation with each position determination. At 24 stations the elevation was measured twice; in 21 cases, the agreement was 8 m or less. Again, the standard deviation of elevation provided by the MX1502 decreased as the number of passes increased in the same manner as for the horizontal position. Seventy-five percent of the differences were within one standard deviation of the least accurate of the two elevations, and no differences were more than two standard deviations of the least accurate elevation.

Optical levelling was also carried out to gain detailed profiles of the ice surface. The distance between the level rod and optical level varied from 150 to 250 m. In early surveys, distances were estimated by Skidoo odometer; some were remeasured using EDM's as part of the strain-rate survey. During later surveys, upper and lower cross hairs were read which gave the distance between level rod positions and also served as a check on errors. Level lines were not closed. On three occasions, optical levelling was carried out between stations which had elevations that were also determined by geoceiver. The difference in elevation determined by each method typically misclosed by 2 to 3 m. This misclosure is fully accounted for by the error in the geoceiver-determined height and is consistent with the given errors. Table 5 gives geographic position of each profile; Table 6 gives the data. Figure 5 shows the geographic locations of each levelling line and Figure 6 presents the elevation profiles for these lines.

TABLE 4: ELEVATIONS FROM GEOCEIVER MEASUREMENTS

Station	Elevation (m)	Number of Passes	Difference (m)	Error (m)	Diff/Error
	YR1 YR2	YR1	YR2		
Crary					
E3	58	45	6	30	1.6
E4	64	66	15	7	
K1	26	20	5	8	
K2	18	18	7	8	
K3	8	4	13	4	
J1	31	37	6	11	
J2	26	31	3	15	
J3	32	29	18	16	
C4	177	212	1	14	
G3	29	33	5	9	
G4	25	33	2	20	
L1	60	59	25	40	
LPL	98	94	19	20	
E2.3A	40	45	17	9	
A2	36	44	10	8	
C2	45	48	58	7	
C3	56	45	47	12	
G2	39	50	12	25	
H2					
Downstream	B				
M0	74	81	12	23	
E19	88	85	5	38	
DNB	91	90	45	23	
G1	114	133	15	14	
G2	135	134	17	11	
A19	100	101	15	20	
B18	84	87	12	23	
B25	103	103	14	23	
H2	125	118	19	38	
M2					
M3	70	56	34	34	
N1	49	52	8	27	
N2	52	47	28	28	
N3	47	70	7	7	
N4					
Downstream	C				
DNC	99	91	25	25	1.2
A1	86	84	24	24	
A2	74	64	15	6	
A3	64	64	15	6	
B2	123	123	??	1	1.6
B3	116	125	16	6	
C1	125	125	8	6	
C3	30	48	4.5	8	
D19	50	50	8	5	
H5					
Wisconsin					
W9	61	22			
W6	50	23			
W5	59	5			
W3	94	20			

TABLE 5: POSITIONS OF OPTICALLY LEVELLED TRAVERSES

Station From->To	Line (Fig. 5)	Start Position LAT	Start Position LONG	End Position LAT	End Position LONG	Dist (km)	Bearing (deg)
		d m s	d m s	d m s	d m s		
CRARY							
A2->H2	2	83 57 37	164 14 14	83 53 17	164 02 45	8.22	15.8
A2->N1	6	83 57 37	164 14 14	83 55 15	163 29 24	10.07	63.8
E2.3A-E2.5	12	83 33 09	168 12 43	83 30 04	168 39 35	8.0	315.3
E2.3A-E2	13	83 33 09	168 12 43	83 36 43	167 40 53	9.0	135.3
E2.2->N	14	83 34 30	168 00 50	83 31 50	167 36 58	7.0	245.3
E2.2->S	15	83 34 30	168 00 50	83 38 38	168 11 07	3.0	225.3
E2.3A->S	16	83 33 09	168 12 43	83 35 14	168 31 38	7.5	225.3
E2.3A->N	17	83 33 09	168 12 43	83 30 29	167 38 56	7.5	245.3
E2.4->S	18	83 31 48	168 24 32	83 31 14	168 19 26	4.4	145.3
E2.4->N	19	83 31 48	172 25 20	83 06 50	172 30 14	4.4	162.4
L1(W)-L1(X)	20	83 06 50	172 24 44	83 06 50	172 44 02	1.5	237.8
L1(X)-L1(Z)	21	83 06 50	172 30 14	83 06 50	172 30 14	1.5	232.8
L1(Z)-L1(D)	22	83 06 50	172 30 14	83 06 50	172 30 14	0.0	225.3
O->G2	23	83 47 14	166 01 28	83 47 33	166 07 59	2.1	226.5
O->D1	24	83 47 14	166 01 28	83 47 33	166 07 59	11.0	19.0
DOWNSTREAM B							
E10-C10	1	84 10 44	155 37 12	84 12 53	153 06 35	40.99	99.3
A19-B10	3	84 24 27	154 42 21	84 02 42	154 07 06	3.92	68.9
N1-N1'	4	83 50 30	161 56 00	83 49 44	161 37 37	40.98	131.6
N1'-N4	5	83 49 44	161 37 37	84 03 44	159 01 17	208.0	208.0
E5->S	7	84 11 30	154 59 30	84 05 04	155 18 29	7.5	27.0
E5->N	8	84 11 30	154 59 30	84 07 19	154 41 32	7.5	27.0
C4->N	9	84 12 31	153 51 48	84 11 19	153 45 32	2.5	207.9
C4->S	10	84 12 31	153 51 48	84 16 20	154 12 07	8.0	212.0
C4-(S16)->S	11	84 16 20	154 12 07	84 17 15	154 17 52	2.0	70.2
G2-G2'	15	84 00 28	150 32 03	83 58 26	153 39 18	13.0	32.1
B10-B18	34	84 02 42	154 07 06	83 57 00	153 14 18	12.62	32.1
DOWNSTREAM C							
O->A1	26	82 49 18	152 26 55	82 54 42	152 27 55	10.0	181.3
A3->O	27	82 49 18	152 26 55	82 16 00	152 31 38	5.0	1.1
O-B3	28	82 49 18	152 26 55	82 18 07	152 23 07	5.5	0.9
O->D19	29	82 49 18	152 26 55	82 09 154	154 25 51	27.5	269.4
D19->O	30	82 49 08	152 25 09	82 09 154	154 25 51	12.0	263.1
B27-H5	31	82 35 26	152 15 43	82 15 43	153 15 43	5.0	326.2
A4-H5	32	82 35 26	152 15 43	82 15 43	153 15 43	5.0	326.2
H5-B2	33	82 34 32	153 15 05	82 34 32	153 19 44	2.0	326.2

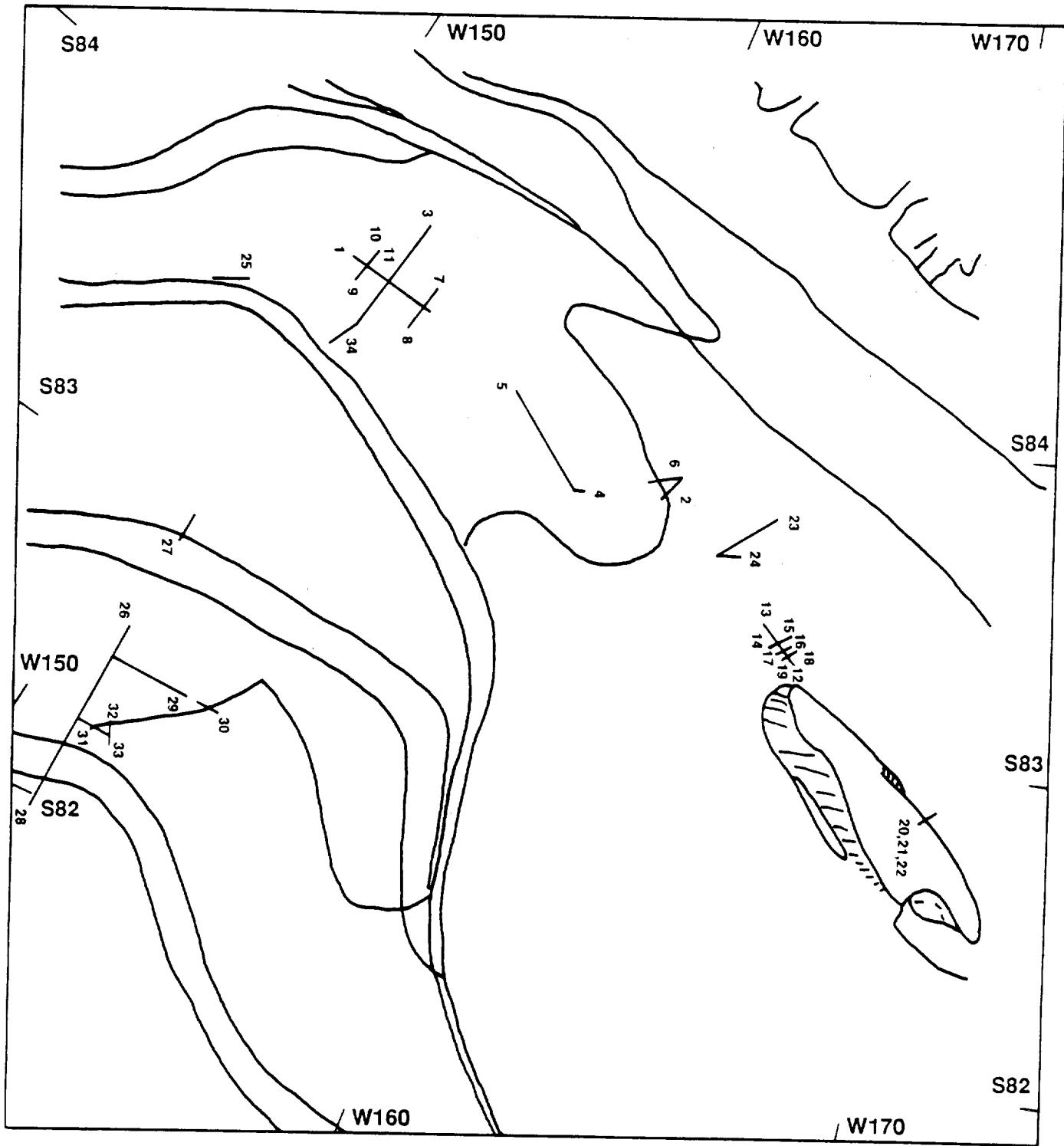


Figure 5: Regional map showing lines of optical levelling.

TABLE 6: SURFACE ELEVATION PROFILES FROM OPTICAL LEVELLING

LINE 1	
Dist (km)	EL (m)
0.00	136.99
0.48	137.38
0.96	138.10
1.44	138.95
1.91	139.66
2.39	140.00
2.86	140.28
3.34	140.30
3.81	140.72
4.29	141.21
4.77	141.69
5.24	142.11
5.72	142.43
6.21	142.81
6.70	142.78
7.19	142.82
7.67	142.85
8.14	142.51
8.62	142.31
9.11	142.08
9.60	142.19
10.09	142.24
10.59	142.27
11.10	142.32
11.60	141.95
12.07	141.56
12.54	141.07
13.01	140.20
13.48	139.62
13.96	139.24
14.43	139.00
15.02	138.91
15.60	139.42
16.08	140.26
16.57	141.50
17.05	142.43
17.53	143.43
18.02	143.50
18.50	143.68
18.98	143.74
19.47	143.65
19.95	143.02
20.43	143.65
20.91	143.68
21.39	142.02
21.88	142.79
22.37	141.80
22.86	141.32
23.33	140.89
23.79	139.97
24.26	138.97
24.72	138.26
25.18	137.34
25.64	136.83
26.13	136.76
26.62	136.73
27.11	136.82
27.57	136.64
28.03	136.32
28.49	136.09

## LINE 2

A2 toward H2

Dist(km)

EL(m)

CIR

0.000	91.82	
0.392	91.82	
0.736	92.27	
1.092	92.16	
1.440	92.05	
1.732	92.05	
2.091	91.72	
2.383	91.14	
2.735	90.77	
3.099	91.02	
3.463	91.34	
3.825	91.00	
4.114	91.08	
4.514	91.19	
4.880	92.34	
5.222	93.76	
5.532	94.66	
5.892	95.63	
6.229	96.13	
6.614	96.28	
7.007	96.03	
7.413	95.99	
7.738	95.60	
8.017	95.75	
8.220	95.77	

LINE 3  
DNB  
A19 to B10

Dist (km)	EL (m)
0.00	150.00
0.50	149.52
1.00	148.01
1.40	147.19
1.90	146.91
2.40	145.57
2.90	145.13
3.40	145.10
3.90	146.53
4.30	146.78
4.80	143.83
5.30	140.49
5.80	138.12
6.30	135.30
6.80	133.88
7.30	133.80
7.70	133.80
8.20	135.25
8.70	135.42
9.20	136.48
9.70	138.79
10.20	140.76
10.70	141.70
11.20	141.39
11.60	139.97
12.10	137.50
12.50	142.23
13.00	143.92
13.50	144.05
14.00	143.65
14.50	144.03
15.00	144.74
15.40	144.74
15.90	146.33
16.40	143.19
16.90	142.85
17.40	144.95
17.80	146.84
18.32	148.08
18.84	146.28
19.34	147.25
20.34	149.23
20.83	150.14
21.29	149.16
21.74	147.59
22.20	145.94
22.67	144.21
23.14	142.09
23.61	139.47
24.10	138.03
24.58	138.80
25.06	139.47
25.51	138.51
25.96	137.32
26.41	135.97
26.86	136.34
27.37	136.97
27.86	137.02
28.38	137.47
29.87	141.59
30.35	143.58
31.32	145.11
32.29	144.44

33	33.27	144.45
33	33.76	144.41
34	34.25	144.09
34	34.74	145.04
35	35.23	146.94
35	35.71	147.89
36	36.20	147.86
36	36.68	146.69
37	37.15	146.23
37	37.62	144.62
38	38.08	142.61
38	38.56	140.66
39	39.03	141.33
39	39.51	143.00
40	40.00	143.61
40	40.49	144.26
40	40.99	143.93

LINE 4  
DNB  
N1 to N1'

Dist (km)	El (m)
0.000	92.37
0.367	93.29
0.746	94.50
1.186	95.29
1.626	96.54
2.056	98.09
2.458	99.93
2.860	101.37
3.260	102.29
3.690	102.46
3.922	102.46

LINE 5  
DNB  
N1' to N4

Dist (km)	El (m)
0.000	102.46
0.347	102.28
0.672	102.24
1.048	102.36
1.410	102.47
1.718	102.94
2.074	103.68
2.449	104.65
2.805	105.03
3.162	105.25
3.597	105.36
4.000	105.27
4.371	105.66
4.732	106.05
5.063	106.50
5.484	107.03
5.819	107.47
6.241	108.07
6.606	108.79
6.986	109.97
7.398	111.10
7.792	111.79
8.224	1112.40
8.606	1112.92
8.982	1112.85
9.360	1112.65
9.729	1112.93
10.132	1113.38
10.538	1113.48
10.933	1113.73
11.320	1114.07
11.724	1114.78
12.099	1115.36
12.546	1115.71
12.926	1115.54
13.377	1115.29
13.782	1115.03
14.208	1115.36
14.600	1115.70
14.981	1116.15
15.427	1116.82
15.809	1117.18
16.285	1119.10
16.665	1120.24
17.102	1121.21
17.512	1122.40
17.937	1123.67
18.345	1124.79
18.780	1125.87
19.160	1126.66
19.598	1127.26

19	959	127	55
20	340	127	55
20	756	127	48
21	139	126	19
21	932	125	14
21	921	123	94
22	288	122	8
22	675	121	78
23	058	120	51
23	769	120	23
24	424	121	08
24	539	120	0
25	906	120	0
25	320	120	0
26	646	120	0
26	787	120	0
27	058	120	0
27	431	120	0
28	735	120	0
28	110	120	0
29	689	120	0
29	580	120	0
29	917	120	0
29	191	120	0
29	848	120	0
29	462	120	0
30	357	120	0
30	102	120	0
31	394	120	0
31	788	120	0
31	040	120	0
31	360	120	0
32	471	120	0
32	173	120	0
32	736	120	0
33	254	120	0
33	890	120	0
34	686	120	0
34	354	120	0
35	930	120	0
35	630	120	0
36	686	120	0
36	256	120	0
37	630	120	0
37	955	120	0
38	748	120	0
38	478	120	0
39	740	120	0
39	451	120	0
40	788	120	0

LINE 6  
CIR  
A2 toward N1

Dist(km)	El(m)
0.000	91.00
0.413	91.90
0.767	92.32
1.065	92.38
1.401	92.50
1.752	92.13
2.101	91.87
2.456	91.57
2.754	91.33
3.094	91.19
3.456	91.62
3.794	92.49
4.142	94.11
4.459	95.27
4.787	96.31
5.135	96.31
5.497	96.12
5.833	97.43
6.193	97.35
6.543	97.45
6.881	97.35
7.122	97.35
7.434	97.35
7.771	97.35
8.125	97.35
8.420	97.35
8.717	97.35
9.046	97.35
9.378	97.35
9.716	97.35
10.071	97.35

LINE 7  
DNB  
E5 to E5(s15)

0.0	142.82
0.5	142.74
1.0	143.06
1.5	143.06
2.0	144.68
2.5	144.12
3.0	143.93
3.5	142.63
4.0	139.90
4.5	139.69
5.0	140.88
5.5	141.54
6.0	141.08
6.5	139.92
7.0	138.54
7.5	138.34
8.0	138.67

LINE 8  
DNB  
E5  
to E5(N15)

0.0	142.82
0.5	143.45
1.0	143.64
1.5	142.91
2.0	142.27
2.5	141.20
3.0	139.41
3.5	138.26
4.0	138.32
4.5	138.89
5.0	138.73
5.5	138.99
6.0	140.57
6.5	143.11
7.0	143.60
7.5	141.98

LINE 9  
DNB  
C4  
to C4(N5)

0.0	143.20
0.5	141.56
1.0	140.52
1.5	139.40
2.0	138.89
2.5	139.36

LINE 10  
DNB  
C4  
to C4(S16)

Dist (km)	El (m)
0.0	143.20
0.5	143.66
1.0	142.79
1.5	141.29
2.0	138.99
2.5	135.00
3.0	134.33
4.0	133.80
4.5	133.74
5.0	134.14
5.5	134.23
6.0	135.22
6.5	138.44
7.0	140.71
7.5	139.91
8.0	139.43

LINE 11  
DNB

Dist (km)	El(m)
0.00	139.43
0.50	139.91
1.0	140.87
1.5	141.04
2.0	139.85

LINE 12  
CIR  
E2. 3A to E2.5

Dist(km)	El(m)
0.00	83
0.5	81.37
1.0	79.83
1.5	79.67
2.0	78.85
2.5	76.31
3.0	74.13
3.25	71.95
3.5	71.14
4.0	71.34
4.5	72.32
5.0	72.96
5.5	73.74
6.0	74.16
6.5	72.64
7.0	72.83
7.5	74.48
8.0	75.57

LINE 13  
CIR  
E2. 3A to E2

Dist(km)	El(m)
0.0	83
0.5	83.33
1.0	82.67
1.5	83.55
2.0	85.09
2.5	86.00
3.0	87.28
3.5	87.21
4.0	86.64
4.5	86.27
5.0	85.54
5.5	85.66
6.0	85.04
6.5	85.35
7.0	84.77
7.5	82.81
8.0	80.45
8.5	80.04
9.0	

LINE 14  
CIR  
E2. 2 to E2. 2(N)

Dist(km)	El(m)
0.0	87.21
0.5	86.53
1.0	85.94
1.5	85.35
2.0	85.59
2.5	85.18
3.0	84.89
3.5	84.86
4.0	84.23
4.5	86.32
5.0	87.06
5.5	86.54
6.0	85.89
6.5	85.43
7.0	81.53

LINE 15  
CIR  
E2.2 to E2.2(S)

Dist(km)	El(m)
0.0	87.21
0.5	86.64
1.0	85.51
1.5	81.81
2.0	82.56
2.5	83.48
3.0	82.96

LINE 16  
CIR

E2.3A to E2.3A(S)

Dist(km)	El (m)
0.0	83
0.5	84.20
1.0	83.33
1.5	82.86
2.0	82.38
2.5	82.28
3.0	81.55
3.5	80.95
4.0	80.10
4.5	78.49
5.0	78.83
5.5	77.72

LINE 17  
CIR

E2.3A to E2.3A(N)

Dist(km)	El(m)
0.0	83
0.5	80.84
1.0	81.40
1.5	83.31
2.0	84.43
2.5	84.74
3.0	85.18
3.5	85.85
4.0	87.48
4.5	87.95
5.0	88.53
5.5	87.75
6.0	85.53
6.5	84.01
7.0	82.68

LINE 18  
CIR

E2.4 to E2.4(S)

Dist(km)	El(m)
0.0	71.14
0.5	72.20
1.0	71.68
1.5	72.13
2.0	73.48
2.5	74.05
3.0	73.53
3.5	74.19
4.0	74.45

LINE 19

E2.4 to E2.4(N)

Dist (km)	El (m)
0.0	71.14
0.5	71.49
1.0	71.89
1.5	70.72

LINE 20

LL(W) to LL(X)

Dist (m)	El (m)
0.00	102
334.30	99.855
450.47	99.083

LINE 21

LL(X) to LL(Z)

Dist (m)	El (m)
0.00	99.083
101.40	97.905
317.50	95.213
512.70	92.562
730.50	89.189
898.60	86.525
1126.40	82.789
1276.60	79.974
1402.30	77.377
1447.70	76.154

LINE 22

LL(Z) to LL(D)

Dist (m)	El (m)
0.00	76.154
86.30	74.191
208.20	71.527
363.10	68.803
479.80	67.033
687.60	65.066
838.10	62.503
1014.80	61.740
1217.10	61.445
1433.10	60.725
1639.10	60.845
1858.10	60.675
2075.10	60.465
2345.10	60.015
2567.10	59.765
2846.10	59.475
3068.10	59.465
3367.10	59.355
3633.10	59.285
3817.40	59.100

LINE 23  
CIR  
0 toward C2

Dist(km)	El(m)
0.00	87
0.25	88.02
0.50	88.46
0.75	88.69
1.00	88.60
1.25	89.26
1.50	89.81
1.75	89.90
2.00	89.85
2.25	90.10
2.50	90.09
2.75	90.03
3.00	90.33
3.25	90.95
3.50	91.19
3.75	91.14
4.00	91.14
4.25	91.54
4.50	91.77
4.75	91.01
5.00	91.19
5.25	91.14
5.50	91.23
5.75	91.62
6.00	92.06
6.25	92.23
6.50	92.62
6.75	93.02
7.00	93.42
7.25	93.82
7.50	94.20
7.75	94.58
8.00	94.92
8.25	95.29
8.50	95.67
8.75	96.04
9.00	96.42
9.25	96.79
9.50	97.16
9.75	97.54
10.00	97.91
10.25	98.29
10.50	98.66
10.75	99.04
11.00	99.42
11.25	99.79
11.50	100.16
11.75	100.54
12.00	100.91
12.25	101.28
12.50	101.65
12.75	102.02
13.00	102.39
13.25	102.76
13.50	103.13
13.75	103.50
14.00	103.87
14.25	104.24
14.50	104.61
14.75	104.98
15.00	105.35
15.25	105.72
15.50	106.09
15.75	106.46
16.00	106.83
16.25	107.19
16.50	107.56
16.75	107.93
17.00	108.29
17.25	108.66

17.50	91.59
17.75	91.94
18.00	92.04
18.25	92.20
18.50	92.05
18.75	92.21
19.00	92.11

LINE 24 CIR 0 toward D1	Dist(km)	El(m)
	0.00	87.87
	0.25	87.56
	0.50	87.73
	0.75	87.85
	1.00	87.70
	1.25	87.50
	1.50	87.25
	1.75	87.00
	2.00	86.75
	2.25	86.50
	2.50	86.25
	2.75	86.00
	3.00	85.75
	3.25	85.50
	3.50	85.25
	3.75	85.00
	4.00	84.75
	4.25	84.50
	4.50	84.25
	4.75	84.00
	5.00	83.75
	5.25	83.50
	5.50	83.25
	5.75	83.00
	6.00	82.75
	6.25	82.50
	6.50	82.25
	6.75	82.00
	7.00	81.75
	7.25	81.50
	7.50	81.25
	7.75	81.00
	8.00	80.75
	8.25	80.50
	8.50	80.25
	8.75	80.00
	9.00	80.75
	9.25	80.50
	9.50	80.25
	9.75	80.00
10.00	88.88	87.87
10.25	88.66	87.66
10.50	88.44	87.44
10.75	88.22	87.22
11.00	88.00	87.00

LINE 25 DNB G2 to G2'	
Dist (km)	El (m)
0.0	178
0.5	178.74
1.0	179.42
1.5	180.00
2.0	179.86
2.5	179.18
3.0	177.99
3.5	177.09
4.0	176.65
4.5	176.17
5.0	175.70
5.5	175.51
6.0	175.20
6.5	175.28
7.0	176.34
7.5	177.77
8.0	179.26
8.5	180.80
9.0	181.94
9.5	183.31
10.0	184.61
10.5	185.69
11.0	186.43
11.5	187.69
12.0	188.71
12.5	190.17
13.0	191.65

LINE 26  
DNC  
0 toward A1

Dist(km)	El(m)
0.00	142.142
0.25	142.11
0.50	141.96
0.75	141.38
1.00	141.06
1.25	140.74
1.50	140.54
1.75	140.19
2.00	139.80
2.25	139.41
2.50	139.36
2.75	139.38
3.00	139.11
3.25	138.85
3.50	138.58
3.75	138.66
4.00	138.88
4.25	138.60
4.50	137.90
4.75	137.81
5.00	137.79
5.25	137.71
5.50	137.53
5.75	137.39
6.00	137.11
6.25	137.05
6.50	136.66
6.75	136.15
7.00	136.66
7.25	135.56
7.50	135.14
7.75	135.13
8.00	135.18
8.25	135.13
8.50	135.13
8.75	135.18
9.00	135.13
9.25	135.42
9.50	135.60
9.75	135.57
10.00	

LINE 27  
DNC  
A3 toward 0

Dist(km)	El(m)
0.0	117.00
0.5	115.27
1.0	114.09
1.5	112.76
2.0	111.71
2.5	110.93
3.0	110.24
3.5	108.78
4.0	107.46
4.5	105.20
5.0	103.37

LINE 28  
DNC  
0 to B3

Dist(km)	El(m)
0.00	142.00
0.25	141.91
0.50	141.76
0.75	141.83
1.00	141.92
1.25	142.21
1.50	142.63
1.75	142.83
2.00	143.56
2.25	145.46
2.50	146.79
2.75	147.13
3.00	146.98
3.25	147.21
3.50	147.72
3.75	148.24
4.00	148.29
4.25	148.52
4.50	148.65
4.75	148.77
5.00	148.80
5.25	148.84
5.50	148.87
5.75	148.90
6.00	148.93
6.25	148.96
6.50	148.99
6.75	149.02
7.00	149.05
7.25	149.08
7.50	149.11
7.75	149.14
8.00	149.17
8.25	149.20
8.50	149.23
8.75	149.26
9.00	149.29
9.25	149.32
9.50	149.35
9.75	149.38
10.00	149.41
10.25	149.44
10.50	149.47
10.75	149.50
11.00	149.53
11.25	149.56
11.50	149.59
11.75	149.62
12.00	149.65
12.25	149.68
12.50	149.71
12.75	149.74
13.00	149.77
13.25	149.80
13.50	149.83
13.75	149.86
14.00	149.89
14.25	149.92
14.50	149.95
14.75	149.98
15.00	150.01
15.25	150.04
15.50	150.07
15.75	150.10
16.00	150.13
16.25	150.16
16.50	150.19
16.75	150.22
17.00	150.25



41.5	114.81
42.0	114.84
42.5	114.44
43.0	114.16
43.5	115.53
44.0	117.92
44.5	120.21
45.0	125.91
45.5	129.51
46.0	132.19
46.5	134.86
47.0	138.06
47.5	141.30
48.0	144.81
48.5	148.23
49.0	151.43
49.5	155.16
50.0	159.13
50.5	162.08
51.0	164.03
51.5	165.55
52.0	166.56
52.5	167.56
53.0	168.30
53.5	168.76
54.0	169.08
54.5	169.31
55.0	168.91
55.5	168.22
56.0	167.84
56.5	167.72
57.0	166.60

LINE 29  
DNC  
0 toward D19

Dist (km)	E1 (m)
0.00	141.75
0.25	141.34
0.50	140.99
0.75	140.45
1.00	140.09
1.25	139.69
1.50	139.33
1.75	139.77
2.00	138.22
2.25	138.77
2.50	138.22
2.75	137.81
3.00	137.09
3.25	136.50
3.50	136.01
3.75	135.51
4.00	135.05
4.25	134.59
4.50	134.13
4.75	133.78
5.00	133.22
5.25	132.78
5.50	132.98
5.75	132.13
6.00	131.18
6.25	130.56
6.50	130.19
6.75	129.60
7.00	128.94
7.25	128.81
7.50	127.44
7.75	126.97
8.00	126.72
8.25	126.33
8.50	126.06
8.75	125.75
9.00	125.49
9.25	125.11
9.50	124.72
9.75	124.28
10.00	123.83
10.25	123.09
10.50	122.48
10.75	121.56
11.00	120.89
11.25	120.45
11.50	119.77
11.75	119.38
12.00	118.35
12.25	117.72
12.50	116.98
12.75	116.30
13.00	116.09
13.25	115.58
13.50	115.32
13.75	115.14
14.00	114.88
14.25	114.56
14.50	114.20
14.75	113.93
15.00	113.57
15.25	112.92
15.50	112.13
15.75	111.02
16.00	109.93
16.25	109.21
16.50	109.09
16.75	109.04
17.00	109.00

LINE 30  
DNC

D19 toward 0

Dist (km)	EL(m)
0.0	73.86
0.5	73.86
1.0	73.04
1.5	71.16
2.0	73.85
2.5	82.94
3.0	89.78
3.5	95.04
4.0	98.01
4.5	100.01
5.0	101.49
5.5	103.57
6.0	104.52
6.5	104.98
7.0	105.45
7.5	105.57
8.0	105.92
8.5	106.27
9.0	106.68
9.5	107.09
10.0	107.30
10.5	107.37
11.0	107.24
11.5	107.14
12.0	107.16

LINE 31  
DNC

B27 to H5

Dist (km)	EL(m)
0.0	136.80
0.5	133.82
1.0	130.94
1.5	127.98
2.0	124.79
2.5	121.14
3.0	117.31
3.5	112.94
4.0	108.17
4.5	106.55
5.0	102.78
5.5	100.44
6.0	96.49
6.5	93.09
7.0	89.22
7.5	85.26
8.0	82.12
8.5	79.45
9.0	79.67
9.5	78.76
10.0	78.42
10.5	78.46
11.0	78.39
11.5	78.12
12.0	78.03

LINE 32  
DNC  
A4(RISP) to H5

Dist (km)	El(m)
0.0	115.84
0.25	112.36
0.5	108.38
0.7	105.18
0.9	101.71
1.0	99.87
1.145	97.03
1.425	93.31
1.656	90.03
1.829	87.36
2.0	85.66
2.5	86.60
3.0	87.11
3.5	86.02
4.0	85.10
4.5	85.26
5.0	85

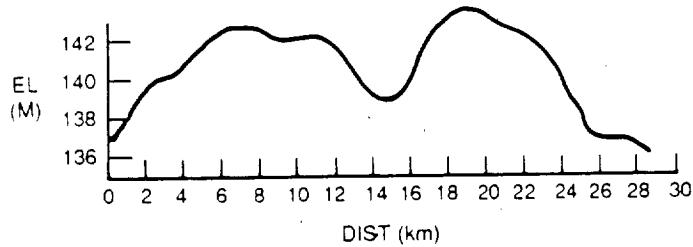
LINE 33  
DNC  
H5-B2(RISP)

Dist (km)	El(m)
0.0	85
0.5	85.30
1.0	86.26
1.5	86.49
2.0	86.52

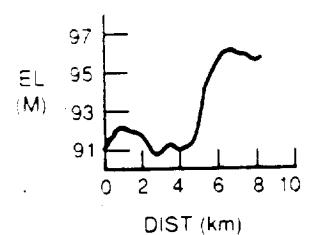
LINE 34  
DNC  
B10-B18

Dist (km)	El(m)
0.0	143.93
0.52	143.79
1.04	143.35
1.55	142.58
2.07	143.17
2.58	144.28
3.09	145.28
3.60	144.58
4.11	143.20
4.62	143.36
5.13	142.48
5.61	141.77
6.09	140.01
6.57	140.28
7.09	141.18
7.61	140.58
8.14	139.14
8.64	138.06
9.14	138.18
9.64	138.27
10.14	138.17
10.63	137.34
11.13	136.03
11.63	136.05
12.12	135.66
12.62	135.35

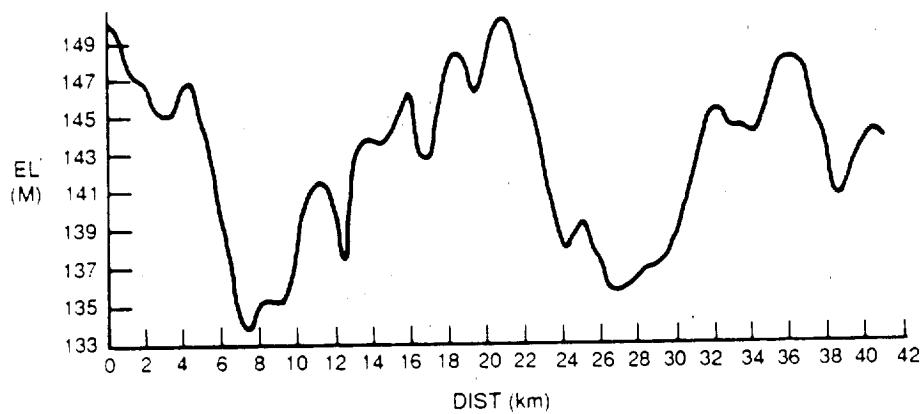
LINE 1  
E10 to C10



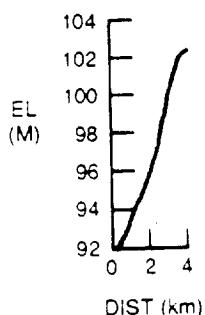
LINE 2  
A2 toward H2



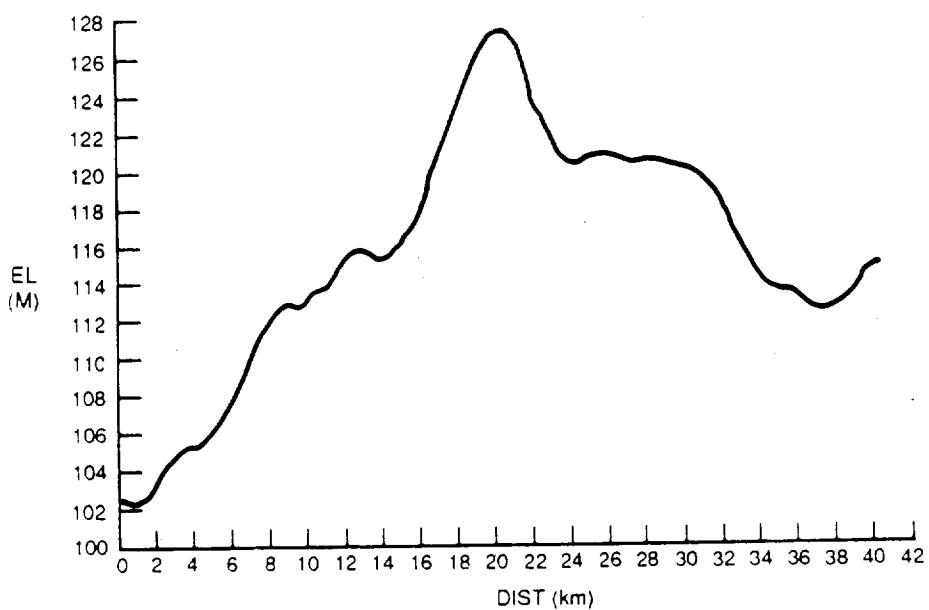
LINE 3  
A19 to B10



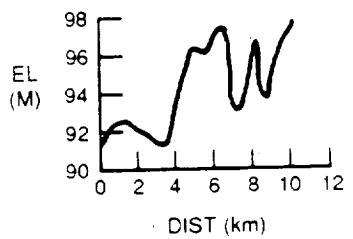
LINE 4  
N1 to N1'



LINE 5  
N1' to N4



LINE 6  
A2 toward N1



LINE 7  
E5 to E5(S15)

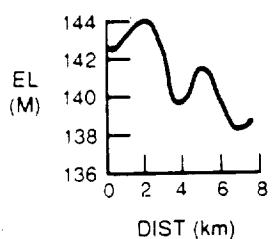
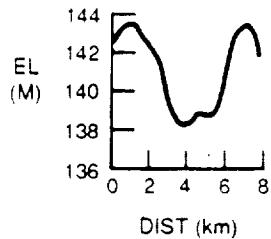
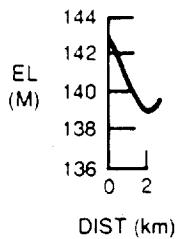


Figure 6: Elevation profiles from optical levelling.

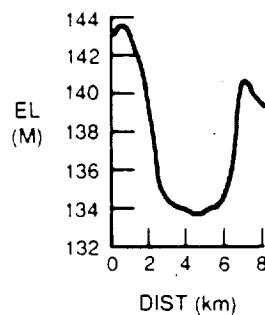
**LINE 8**  
E5 to E5(N15)



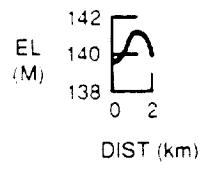
**LINE 9**  
C4 to C4(N5)



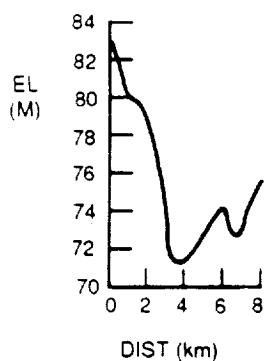
**LINE 10**  
C4 to C4(S16)



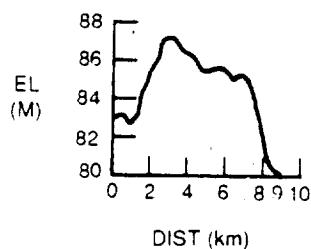
**LINE 11**  
C4(S16) to C4(S20)



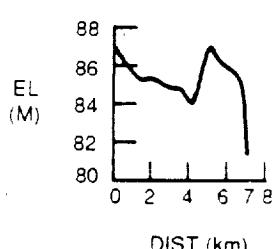
**LINE 12**  
E2.3A to E2.5



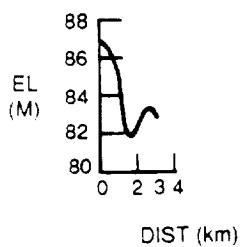
**LINE 13**  
E2.3A to E2



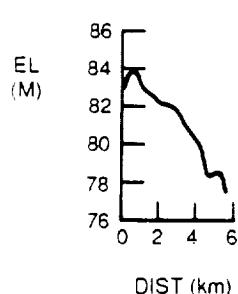
**LINE 14**  
E2.2 to E2.2(N)



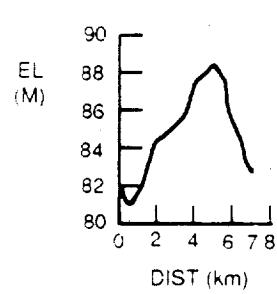
**LINE 15**  
E2.2 to E2.2(S)



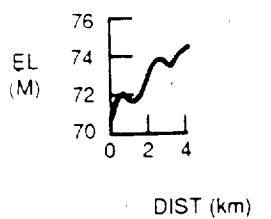
**LINE 16**  
E2.3A to E2.3A(S)



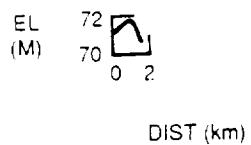
**LINE 17**  
E2.3A to E2.3A(N)



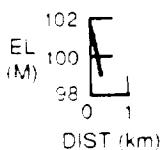
**LINE 18**  
**E2.4 to E2.4(S)**



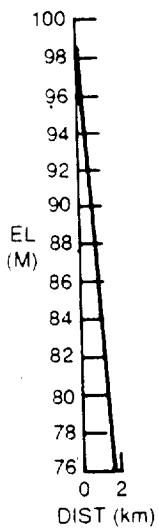
**LINE 19**  
**E2.4 to E2.4(N)**



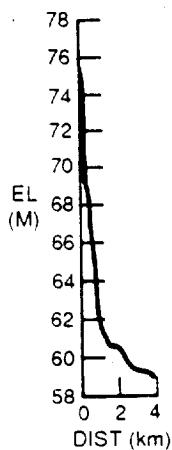
**LINE 20**  
**L1(W) to L1(X)**



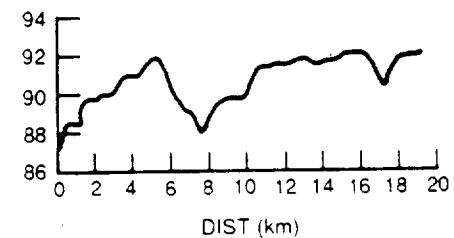
**LINE 21**  
**L1(X) to L1(Z)**



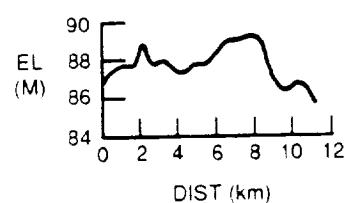
**LINE 22**  
**L1(Z) to L1(D)**



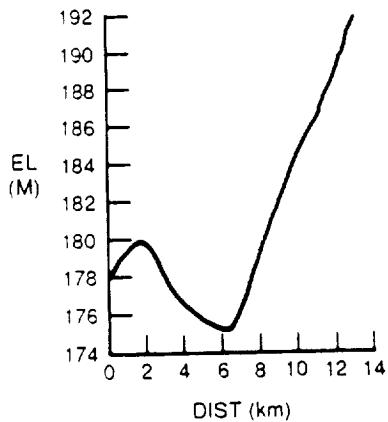
**LINE 23**  
**0 toward C2**



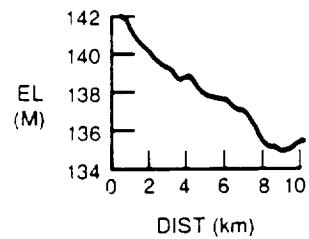
**LINE 24**  
**0 toward D1**



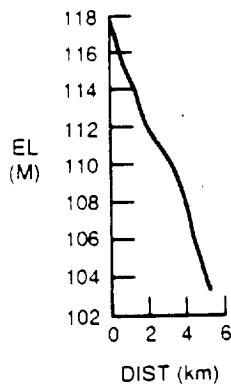
**LINE 25**  
**G2 to G2'**



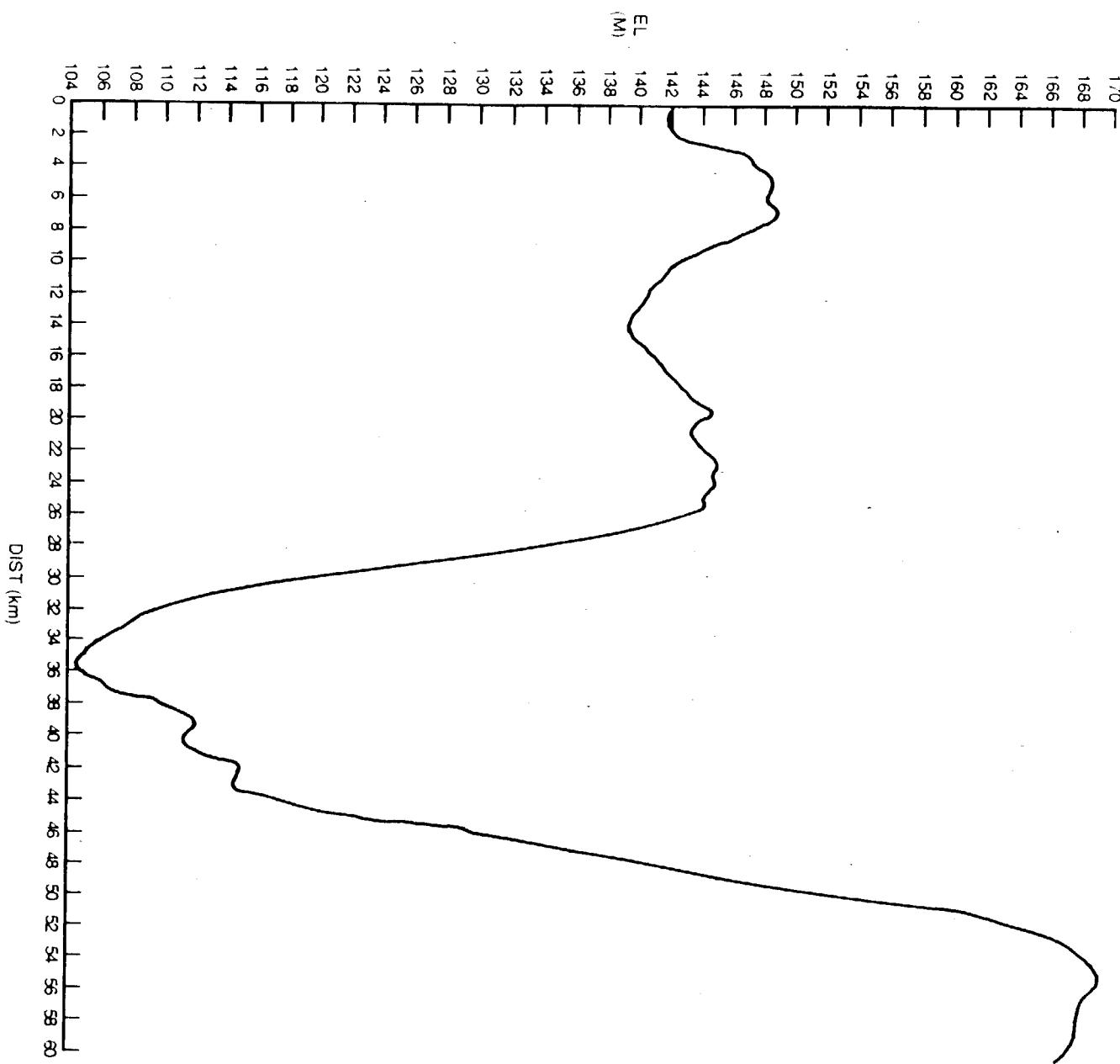
**LINE 26**  
**0 toward A1**



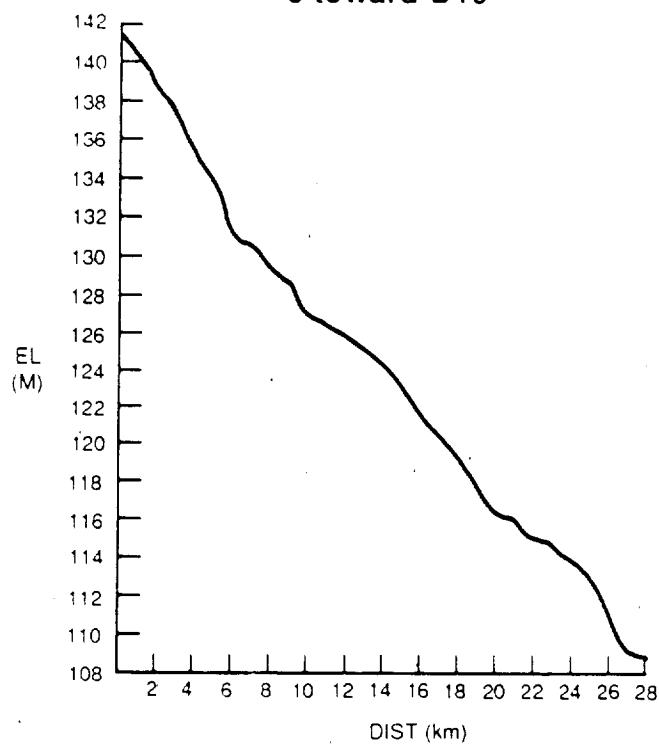
**LINE 27**  
**A3 toward 0**



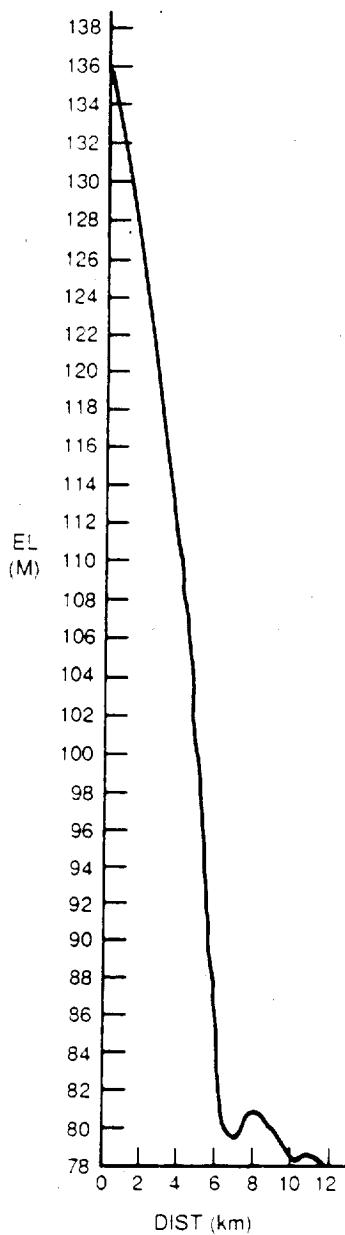
LINE 28  
0 to B3



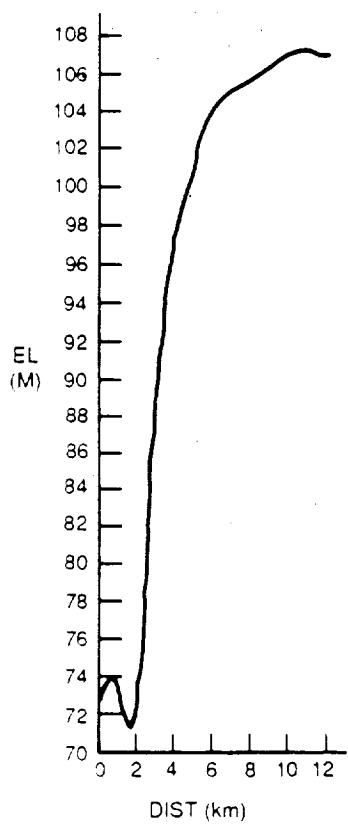
LINE 29  
0 toward D19



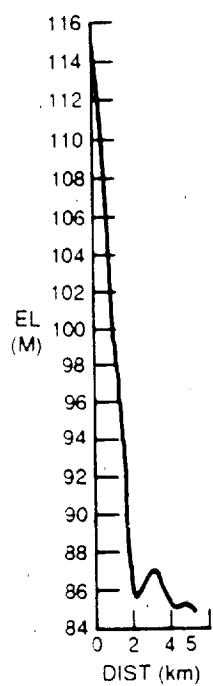
LINE 31  
B27 to H5



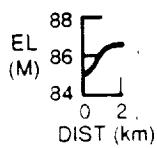
LINE 30  
D19 toward 0



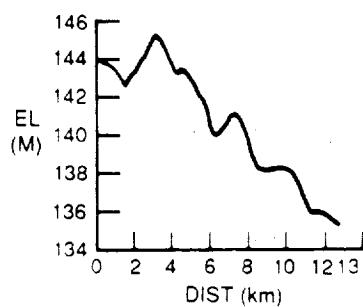
**LINE 32**  
A4(RISP) to H5



**LINE 33**  
H5 to B2(RISP)



**LINE 34**  
B10 to B18



#### SECTION 4: RADIO-ECHO SOUNDING OF ICE THICKNESS

Spot measurements of ice thickness were made at several survey stations (Table 7) and along three short profiles near the grounding line of Ice Stream B (Tables 8, 9 and 10 and Figure 7). The radio-echo sounder used was on loan from the University of Wisconsin, Geophysics and Polar Science Center. A description of this radio-echo system is given by Shabtaie and Bentley (1987). To reduce the two-way travel time to ice thickness, a velocity of  $169 \times 10^{-6}$  m/s was used for the radar energy (Shabtaie, personal communication).

TABLE 7: STATION ICE THICKNESS

STATION	TWO-WAY TRAVEL TIME	THICKNESS (m)
	( $\mu$ s)	
M2	8.27	695
M3	8.83	742
N1	7.67	649
N2	6.83	580
N3	7.56	638
N4	8.62	728
C2	8.16	690
H2	7.65	646
A2	8.18	691

TABLE 8: THICKNESS PROFILE FROM N1 TO N4

STATION	TWO-WAY TRAVEL TIME	THICKNESS (m)
(km)	( $\mu$ s)	
0	7.68	649
1.0	7.63	645
2.0	7.71	651
3.0	7.74	654
4.0	7.84	658
5.0	7.90	662
6.0	7.92	668
7.0	7.98	669
8.0	8.02	674
9.0	8.01	677
10.0	8.27	684
11.0	8.34	695
12.0	8.39	698
13.0	8.41	692
14.0	8.44	699
15.0	8.51	705
16.0	8.54	709
17.0	8.59	713
18.0	8.61	719
19.0	8.64	723
20.0	8.67	727
21.0	8.70	728
22.0	8.73	728
23.0	8.76	730
24.0	8.79	731
25.0	8.82	732
26.0	8.85	732
27.0	8.88	730
28.0	8.91	731
29.0	8.94	732
30.0	8.97	733
31.0	9.00	734
32.0	9.03	
33.0	9.06	
34.0	9.09	
35.0	9.12	
36.0	9.15	
37.0	9.18	
38.0	9.21	
39.0	9.24	
40.0	9.27	
41.0	9.30	
42.0	9.33	
43.0	9.36	
44.0	9.39	
45.0	9.42	
46.0	9.45	
47.0	9.48	
48.0	9.51	
49.0	9.54	
50.0	9.57	
51.0	9.60	
52.0	9.63	
53.0	9.66	
54.0	9.69	
55.0	9.72	
56.0	9.75	
57.0	9.78	
58.0	9.81	
59.0	9.84	
60.0	9.87	
61.0	9.90	
62.0	9.93	
63.0	9.96	
64.0	10.00	
65.0	10.03	
66.0	10.06	
67.0	10.09	
68.0	10.12	
69.0	10.15	
70.0	10.18	
71.0	10.21	
72.0	10.24	
73.0	10.27	
74.0	10.30	
75.0	10.33	
76.0	10.36	
77.0	10.39	
78.0	10.42	
79.0	10.45	
80.0	10.48	
81.0	10.51	
82.0	10.54	
83.0	10.57	
84.0	10.60	
85.0	10.63	
86.0	10.66	
87.0	10.69	
88.0	10.72	
89.0	10.75	
90.0	10.78	
91.0	10.81	
92.0	10.84	
93.0	10.87	
94.0	10.90	
95.0	10.93	
96.0	10.96	
97.0	11.00	
98.0	11.03	
99.0	11.06	
100.0	11.09	

TABLE 9: THICKNESS PROFILE FROM A2 (CIR) TOWARD N1

STATION (km)	TWO-WAY TRAVEL TIME ( $\mu$ s)	THICKNESS (m)
0.0	8.21	694
1.0	8.28	700
1.5	8.29	701
2.0	8.31	702
2.5	8.27	703
3.0	8.25	699
3.2	8.25	698
4.0	8.27	697
4.4	8.27	699
5.0	8.30	701
5.5	8.28	699
6.0	8.25	697
7.0	8.27	699
7.2	8.26	694
8.0	8.20	693
8.2	8.22	694
9.0	8.24	696
9.2	8.23	696
10.0		

TABLE 10: THICKNESS PROFILE FROM A2 (CIR) TOWARD H2 (CIR)

STATION (km)	TWO-WAY TRAVEL TIME ( $\mu$ s)	THICKNESS (m)
0.0	8.24	696
0.5	8.31	702
1.0	8.32	703
1.5	8.27	698
2.0	8.24	695
2.2	8.26	698
2.3	8.20	692
2.4	8.24	691
2.6	8.18	692
2.7	8.17	690
3.0	8.18	697
3.2	8.16	690
3.5	8.15	689
4.0	8.11	685

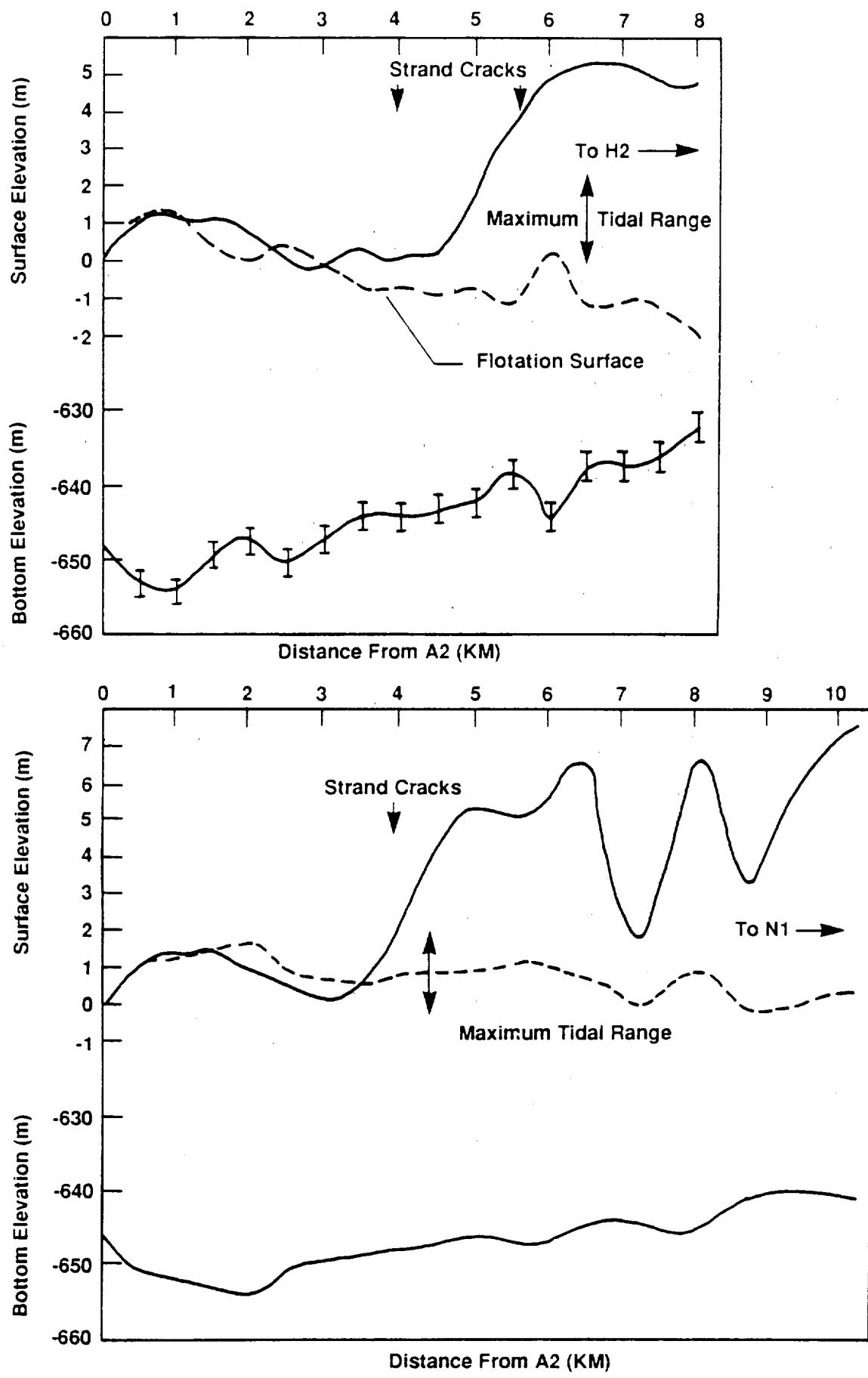


Figure 7: Base and surface topography along two lines near Crary station A2.

## SECTION 5: ACCUMULATION RATE

Two methods were used to measure the accumulation rate. The preferred method involved coring to a depth of about 10 m and sampling the core for beta particle activity. There are peaks in this measurement in the 1956 and 1965 horizons associated with global nuclear bomb tests. Cores have been taken at four sites, (DNB, CIR, E19 and DNC) but the analysis has not been completed.

The second method is to repeat measurements of the length of exposed survey stakes. The normal time interval between measurements of stake heights was only 1 year, so the measurement of accumulation rate may not be representative of the longer term average. To convert these stake exposure measurements to mass equivalents, we suggest a value of  $388 \text{ kg/m}^{-3}$  as the average density of the surface firn.

Table 11 includes these measurements indicating location and the time interval. The standard deviation of the accumulation rate is derived from the eight stake exposures (four survey poles and four adjacent bamboo poles) measured at each strain rosette covering an area of typically  $7 \text{ km}^2$ . The stake lines at Downstream B usually consist of 10 stakes along an 15-km line.

TABLE 11: ACCUMULATION RATES FROM STAKE EXPOSURES AT STRAIN ROSETTES

STATION	FIRN ICE*			STATION	FIRN ICE*		
	ACCUM (cm/a)	EQUIV (cm/a)	DATE		ACCUM (cm/a)	EQUIV (cm/a)	DATE
<b>Crary</b>							
A1	50	11	21	5	83,84		
A2	58	13	24	6	83,84		
	24	7	10	3	84,85		
B1	50	6	21	2	83,84		
B2	70	11	30	5	83,84		
C1	70	4	30	2	83,84		
C2	68	12	29	5	83,84		
	29	8	12	3	84,85		
C3	41	12	17	5	84,85		
C4	42	11	18	5	84,85		
D1	62	14	26	6	83,84		
D2	65	4	28	2	83,84		
E'1	20	3	8	1	84,85		
E'2	20	12	8	5	84,85		
E1	62	3	26	1	84,85		
E2	53	11	22	5	84,85		
E2.3	19	2	8	1	84,85		
E2.5	24	5	10	2	84,85		
E3	30	13	13	6	84,85		
E4	31	17	13	7	84,85		
F1	53	18	22	8	83,84		
F2	60	15	25	6	83,84		
G1	51	12	22	5	83,84		
G2	49	9	21	4	83,84		
G3	16	5	7	2	84,85		
G4	16	14	7	6	84,85		
H1	54	5	23	2	83,84		
H2	48	1	20	1	83,84		
J1	32	28	14	12	84,85		
J2	29	2	12	1	84,85		
J3	31	5	13	2	84,85		
K1	21	8	9	3	84,85		
K2	20	6	8	2	84,85		
K3	7	4	3	2	84,85		
L1	30	12	13	5	84,85		
O	20	5	8	2	84,85		
<b>Downstream B</b>							
All	45	3	19	1	83,84		
B25	11	3	5	1	84,85		
G1	15	7	6	3	84,85		
G2	13	5	6	2	84,85		
M0	22	2	9	1	84,85		
M1	27	9	11	4	84,85		
M2	26	3	11	1	84,85		
M3	24	9	10	4	84,85		
M4	29	11	12	5	84,85		
M5	19	3	8	1	84,85		
M6	23	6	10	2	84,85		

\*Density of firn= 388 kg m<sup>-3</sup>ORIGINAL PAGE IS  
OF POOR QUALITY

## SECTION 6: TEN-METER TEMPERATURE AND FIRN DENSITY PROFILES

At the locations where 10-meter-long cores were removed, temperature measurements were made in the resulting holes. In 1983, a single thermistor was used at a single depth, while in 1984 and 1985 a thermistor chain was hung in the hole to measure the vertical temperature gradient. The single thermistors were glass bead Fenwall GB32M2 (2000-ohm resistance at 25 degrees C). The thermistor chain alternated these with Fenwall GB41P2 thermistors (10,000-ohm resistance at 25 degrees C) every meter for 5 m. Each thermistor was calibrated at zero and three sub-zero temperatures. A precision Wheatston bridge (Leeds and Northrup Model 4289-3) was used. The hole was covered to reduce air circulation. Repeated measurements were made as soon as the core was removed to allow the cooling curve to be determined, giving a more accurate estimate of the equilibrium temperature. Measurements were continued until the cooling rate was less than 0.01 degrees C per hour, which usually occurred after 4 to 6 hours. When possible, a final measurement was made after many hours to confirm the calculated equilibrium temperature. Self-heating of the thermistors caused by the 1-mA current output by the measurement bridge limited the accuracy of these temperature measurements to  $\pm 0.05$  degrees C. Table 12 presents these data.

Firn density profiles were calculated using volume and weight data collected at sites where 10-meter-long cores were extracted (Figure 8 and Table 13). In the process of preparing the 10-meter core for shipment, each core was cut into 10-cm sections and weighed to the nearest tenth of a gram using a sheltered triple-beam balance. During the 1983 and 1984 field seasons the weighing was done after each 10-cm section had been placed in a small plastic bag, while in 1985 the core sections were weighed beforehand. The weight of the bags is considered negligible and within the error of these measurements; hence no attempt is made here to adjust the 1983 and 1984 data. In 1983 and 1984 the diameter of the core was taken to be that of the inside diameter of the coring assembly (7.3 cm), while in the 1985 season, each

core section was measured individually using a caliper. The 1985 data showed that the core's diameter varied from the inside diameter of the coring assembly by less than 2mm. Density values obtained from the 10-cm segments of core are assigned to the vertical position of the center point of each segment.

TABLE 12: TEMPERATURE MEASUREMENTS

Station	Date			Depth (m)	Temperature (C)
	M	/	D/Y		
CIR	12	6	83	10.2	-25.7
	12	2	84	4.0	-26.1
				5.0	-25.5
				6.0	-25.8
				7.0	-25.2
				8.0	-25.1
				9.0	-25.1
				10.0	-25.4
	12	3	84	0.0	-12.0
				1.0	-19.2
				2.0	-24.3
				3.0	-26.0
				4.0	-25.8
				5.0	-26.7
				6.0	-26.7
DNB	11	17	83	9.5	-26.2
	11	19	83	9.5	-26.2
	12	15	84	3.7	-25.9
				4.7	-25.7
				5.7	-26.3
				6.7	-25.2
				7.7	-25.5
				8.7	-25.6
				9.7	-26.1
E19	12	24	85	8.0	-25.9
				9.0	-25.3
				10.0	-26.2
DNC	12	14	83	9.6	-29.8

TABLE 13 (a): FIRN DENSITY PROFILES  
 DNB CAMP  
 DEC 1984

DEPTH (m)	DENSITY (gm cm <sup>-3</sup> )	DEPTH (m)	DENSITY (gm cm <sup>-3</sup> )
0.00	0.33	5.00	0.402
0.23	0.42	5.00	0.460
0.33	0.44	5.00	0.434
0.42	0.44	5.00	0.450
0.50	0.44	5.00	0.437
0.57	0.44	5.00	0.448
0.63	0.44	5.00	0.452
0.69	0.44	5.00	0.450
0.76	0.44	5.00	0.457
0.85	0.44	5.00	0.464
0.95	0.44	5.00	0.471
1.05	0.44	5.00	0.478
1.14	0.44	5.00	0.484
1.24	0.44	5.00	0.491
1.34	0.44	5.00	0.498
1.44	0.44	5.00	0.504
1.54	0.44	5.00	0.509
1.64	0.44	5.00	0.514
1.74	0.44	5.00	0.519
1.84	0.44	5.00	0.524
1.94	0.44	5.00	0.529
2.04	0.44	5.00	0.534
2.14	0.44	5.00	0.539
2.24	0.44	5.00	0.544
2.34	0.44	5.00	0.549
2.44	0.44	5.00	0.554
2.54	0.44	5.00	0.559
2.64	0.44	5.00	0.564
2.74	0.44	5.00	0.569
2.84	0.44	5.00	0.574
2.94	0.44	5.00	0.579
3.04	0.44	5.00	0.584
3.14	0.44	5.00	0.589
3.24	0.44	5.00	0.594
3.34	0.44	5.00	0.599
3.44	0.44	5.00	0.604
3.54	0.44	5.00	0.609
3.64	0.44	5.00	0.614
3.74	0.44	5.00	0.619
3.84	0.44	5.00	0.624
3.94	0.44	5.00	0.629
4.04	0.44	5.00	0.634
4.14	0.44	5.00	0.639
4.24	0.44	5.00	0.644
4.34	0.44	5.00	0.649
4.44	0.44	5.00	0.654
4.54	0.44	5.00	0.659
4.64	0.44	5.00	0.664
4.74	0.44	5.00	0.669
4.84	0.44	5.00	0.674
4.94	0.44	5.00	0.679
5.04	0.44	5.00	0.684

TABLE 13(b) : FIRN DENSITY PROFILES

DEPTH (m)	DENSITY (gm cm <sup>-3</sup> )	DEPTH (m)	DENSITY (gm cm <sup>-3</sup> )	DEPTH (m)	DENSITY (gm cm <sup>-3</sup> )
114	1.024	113	1.024	112	1.024
111	1.024	110	1.024	109	1.024
108	1.024	107	1.024	106	1.024
105	1.024	104	1.024	103	1.024
102	1.024	101	1.024	100	1.024
98	1.024	97	1.024	96	1.024
94	1.024	93	1.024	92	1.024
90	1.024	89	1.024	88	1.024
85	1.024	84	1.024	83	1.024
80	1.024	79	1.024	78	1.024
75	1.024	74	1.024	73	1.024
70	1.024	69	1.024	68	1.024
65	1.024	64	1.024	63	1.024
60	1.024	59	1.024	58	1.024
55	1.024	54	1.024	53	1.024
50	1.024	49	1.024	48	1.024
45	1.024	44	1.024	43	1.024
40	1.024	39	1.024	38	1.024
35	1.024	34	1.024	33	1.024
30	1.024	29	1.024	28	1.024
25	1.024	24	1.024	23	1.024
20	1.024	19	1.024	18	1.024
15	1.024	14	1.024	13	1.024
10	1.024	9	1.024	8	1.024
5	1.024	4	1.024	3	1.024
0	1.024				

TABLE 13(c) : FIRN DENSITY PROFILES

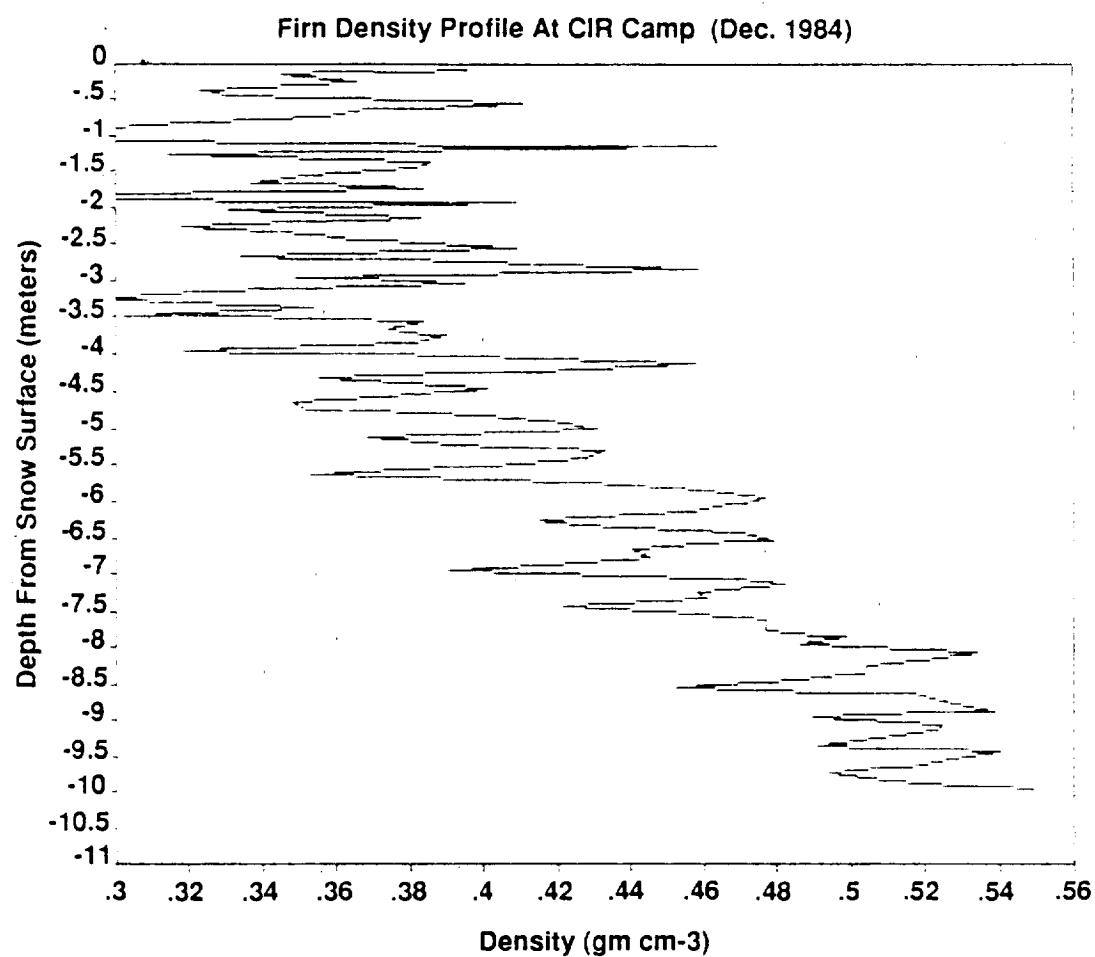
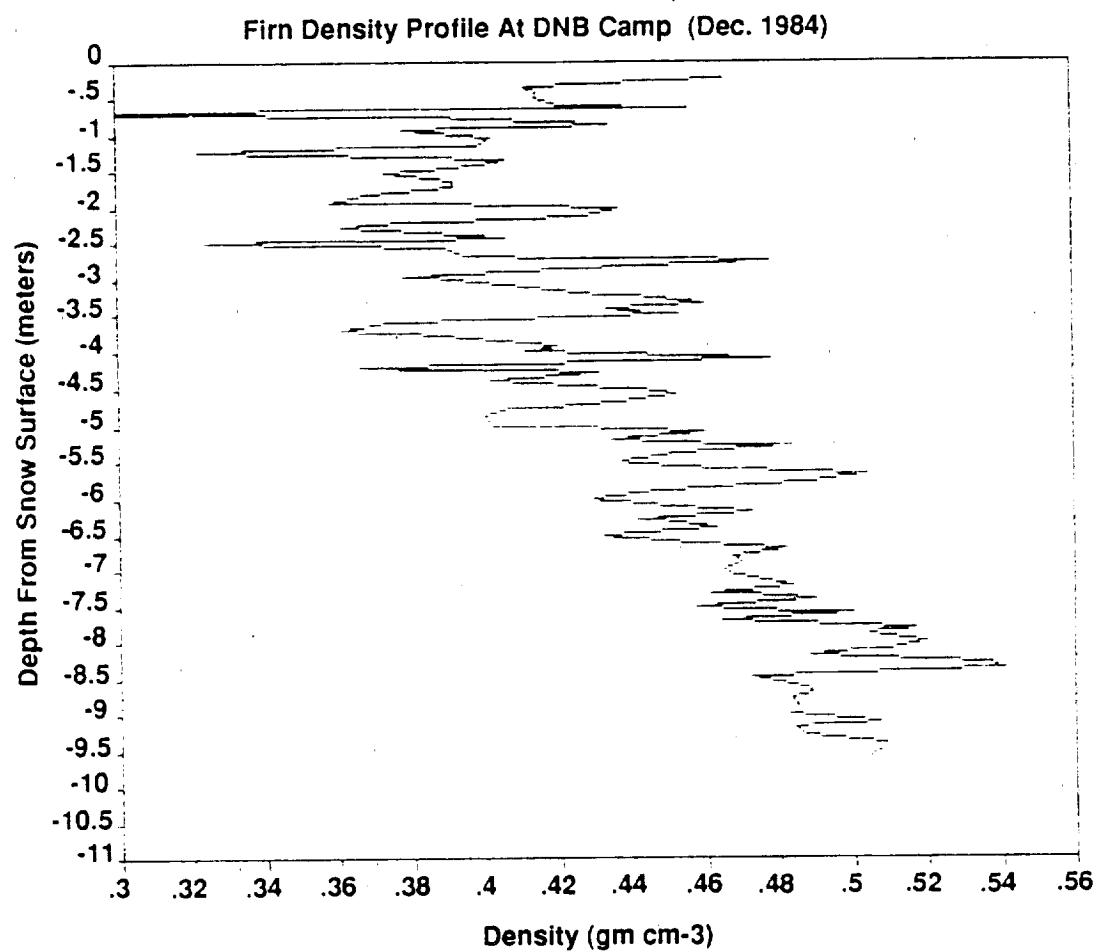
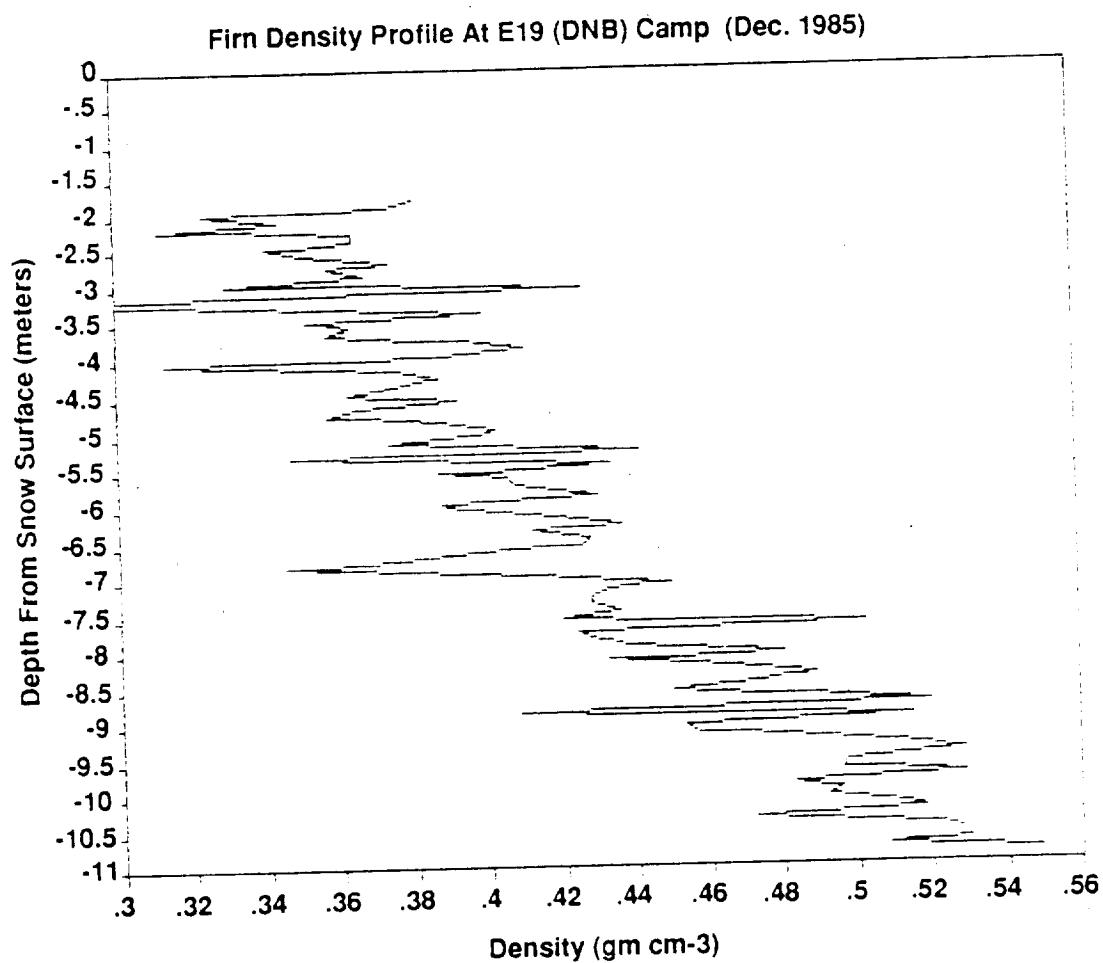


Figure 8 (a-c): Firn density profiles.





**SECTION 7: SPECIAL SITES**  
**RESULTS FROM DOUBLE STAKE LINES AT L1 AND DNB**

Special sites refer to two areas where double stake lines were established to determine the detailed strain field in an extended region. The relative velocities and strain rate were determined as previously described in Section 2

TABLE 14: VELOCITY PROFILE FROM DOUBLE LINE OF STAKES WHICH CROSSES  
THE EDGE OF CRARY ICE RISE NEAR STATION L1 (see Fig 11(a))

STATION NAME	COORDINATES*		VELOCITY	
	X (m)	Y (m)	X-comp. (m/a)	Y-comp. (m/a)
W	50000.00	50000.00	0.00	0.00
X	49590.38	50187.44	0.51	1.06
A3	50468.18	50180.27	0.52	-1.13
Y	49741.70	50964.61	2.75	1.02
A2	50505.86	50981.21	2.73	-0.88
Z	49827.13	51614.92	4.78	1.22
A1	50536.26	51583.82	4.67	-0.63
A0	49860.28	52300.72	9.66	1.46
O	50572.51	52265.96	8.56	-0.41
A	49879.28	52737.89	22.90	2.33
1	50782.65	52820.14	23.85	-0.02
B	49886.65	53322.97	42.55	3.56
2	50752.39	53544.42	48.81	1.67
C	49914.35	54282.94	74.44	5.04
3	50747.48	54509.57	81.04	2.21
D	49954.55	55430.03	109.08	5.00
4	50852.28	55476.13	110.58	2.31
E	49991.85	56612.34	141.96	4.56
5	50989.94	56690.43	144.82	1.13
F	50027.91	57774.16	173.96	3.52
6	51043.10	57828.09	176.34	0.46
G	50062.47	58725.45	197.69	2.64
7	51085.83	58867.63	202.36	-0.56
H	50093.00	59677.40	219.57	1.96
8	51135.38	59928.41	226.66	-1.83
I	50125.66	60685.17	241.91	1.02
9	51175.01	60694.78	243.77	-2.77
J	50190.23	61808.99	265.53	-0.33
10	51231.68	61743.33	265.92	-4.23

\*Azimuth of X-axis is 316.9 degrees. Geographic coordinates of station W are: S83°06'11", W172° 25'23".

TABLE 15: STRAIN RATE PROFILE FROM DOUBLE LINE OF STAKES WHICH CROSSES  
THE EDGE OF CRARY ICE RISE AT STATION L1 (see Fig. 9)

STATIONS	COORDINATES*	$P_1^+$ ( $10^{-3} \text{a}^{-1}$ )	$P_2^+$ ( $10^{-3} \text{a}^{-1}$ )	$P_3^+$ ( $10^{-3} \text{a}^{-1}$ )	ANGLE OF $P_1^+$ WRT X-axis
X(m)	Y(m)				
W X A3	50019 49933	50122 50444	$0.2 \pm 0.2$ 0.52 0.06	$-0.0 \pm 0.2$ -0.04 0.09	$-0.2 \pm 0.3$ -0.5 0.1
X A3 Y	50238	50708	0.45 0.06	-0.12 0.09	76.70 0.08
A3 Y A2	50025	51187	0.75 0.07	-0.2 0.1	69.65 0.05
Y A2 Z	50290	51393	0.69 0.08	-0.1 0.1	65.18 0.07
Z A1 A0	50074	51833	2.59 0.08	-1.95 0.08	46.99 0.01
A1 A0 O	50323	52050	1.39 0.09	-2.21 0.08	59.48 0.01
AO O A	50104	52434	14.9 0.1	-12.9 0.1	47.24 0.00
O A 1	50411	52608	12.92 0.08	-12.69 0.08	48.65 0.00
A 1 B	50182	52960	15.61 0.08	-15.49 0.08	48.83 0.00
1 B 2	50473	53229	16.26 0.07	-15.60 0.07	48.41 0.00
B 2 C	50184	53716	15.57 0.06	-15.23 0.06	47.70 0.00
2 C 3	50471	54112	14.63 0.06	-15.26 0.06	46.64 0.00
C 3 D	50205	54704	13.30 0.06	-13.50 0.05	45.42 0.00
3 D 4	50518	55138	14.04 0.06	-13.50 0.06	45.35 0.00
D 4 E	50266	55839	12.39 0.05	-12.43 0.05	44.40 0.00
4 E 5	50611	56259	12.42 0.05	-12.34 0.05	43.56 0.00
E 5 F	50336	57025	12.05 0.05	-12.14 0.05	43.23 0.00
F 5 G	50687	57430	12.57 0.05	-12.14 0.05	43.45 0.00
F 6 G	50377	58109	11.12 0.05	-10.90 0.06	42.60 0.00
6 G 7	50730	58474	11.14 0.05	-10.91 0.05	42.46 0.00
G 7 H	50413	59090	10.37 0.06	-9.62 0.06	42.15 0.00
H 8 I	50771	59491	9.92 0.05	-9.67 0.05	41.59 0.00
H 8 I	50451	60097	9.72 0.05	-9.09 0.05	41.50 0.00
I 9 J	50812	60436	9.68 0.06	-9.15 0.07	41.03 0.00
I 9 J	50497	61063	9.05 0.06	-8.47 0.07	40.78 0.00
J 10	50865	61416	8.98 0.07	-8.47 0.06	40.23 0.00

\*Azimuth of X-axis is 316.9 degrees

<sup>†</sup> $P_1$  and  $P_2$  are principal axes of horizontal tension and compression, respectively.  $P_3 = -(P_1 + P_2)$  is vertical strain rate

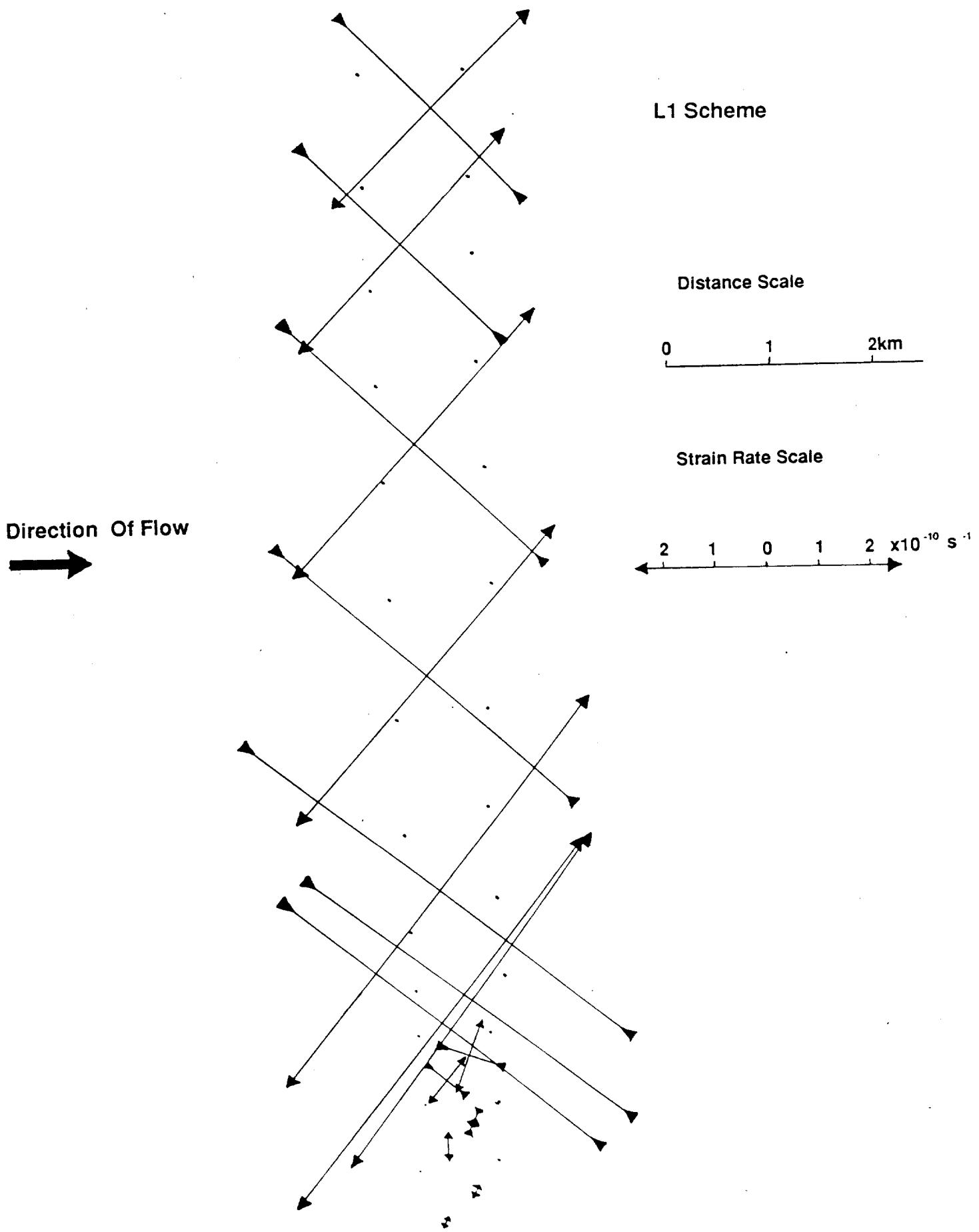


Figure 9: Strain rates measured along a double stake line at L1, across the margin of Crary Ice Rise.

TABLE 16: VELOCITY PROFILE FROM DOUBLE LINE OF STAKES PARALLEL  
TO FLOW IN THE MOUTH OF ICE STREAM B.

STATION NAME	COORDINATES*		VELOCITY	
	X (m)	Y (m)	X-comp. (m/a)	Y-comp. (m/a)
F10	48722.25	65058.83	-47.29	495.84
E10	47123.69	64442.66	-53.71	498.97
F9	48667.88	63612.12	-51.58	497.55
E9	47083.31	63002.17	-57.15	500.27
F8	48614.64	62199.78	-55.19	498.63
E8	47043.61	61580.14	-60.07	501.56
F7	48559.75	60746.05	-58.74	499.84
E7	47004.03	60149.06	-63.26	502.21
F6	48504.94	59312.11	-61.99	501.06
E6	46965.05	58715.93	-65.90	502.75
F5	48450.11	57894.97	-64.67	502.02
E5	46925.04	57248.23	-68.72	503.54
F4	48394.46	56468.13	-67.50	502.78
E4	46885.73	55817.89	-70.98	504.50
F3	48338.80	55037.41	-70.06	504.54
E3	46845.75	54354.00	-74.02	505.78
F2	48283.48	53615.69	-73.13	506.66
E2	46804.46	52842.80	-78.67	507.22
F1	48238.29	52169.61	-77.17	508.23
E1	46765.76	51429.91	-83.15	508.76
A1	48660.00	50000.00	-83.96	510.92
A2	47208.94	50000.00	-86.21	510.74
D1	48138.31	49481.45	-86.68	511.58
C1	46786.96	48832.68	-91.37	511.56
D2	48091.78	48221.00	-91.18	512.99
C2	46861.04	47383.77	-96.03	513.81
D3	48040.47	46856.76	-95.49	514.64
C3	46936.43	45934.86	-100.61	515.35
D4	47991.46	45543.27	-99.90	516.02
C4	47013.62	44481.85	-104.74	516.90
D5	47942.21	44218.70	-104.21	517.35
C5	47090.69	43038.58	-108.88	518.34
D6	47892.79	42887.81	-108.17	518.78
C6	47169.13	41569.26	-113.30	519.28
D7	47843.64	41564.22	-112.25	519.62
C7	47243.96	40177.83	-118.18	520.49
D8	47778.92	39804.28	-118.64	521.08
C8	47318.40	38791.39	-121.23	521.17
D9	47730.55	38479.92	-121.70	521.77
C9	47396.86	37328.70	-125.18	522.68
D10	47683.85	37153.12	-125.62	523.23
C10	47471.54	35941.58	-128.17	523.51

\*Azimuth of X-axis is 289.7 degrees. Geographic coordinates of station A1 are: S84° 10'52", W154° 10'08".

TABLE 17: STRAIN RATE PROFILE FROM DOUBLE LINE OF STAKES WHICH PARALLELS ICE FLOW IN THE MOUTH OF ICE STREAM B (see Fig. 10)

STATIONS	COORDINATES*		$P_1$ ( $10^{-3} \text{a}^{-1}$ )	$P_2$ ( $10^{-3} \text{a}^{-1}$ )	$P_3$ ( $10^{-3} \text{a}^{-1}$ )	ANGLE OF $P_1$ WRT X-axis*
	X(m)	Y(m)				
F10 E10 F9	48171	64371	3.02 ± 0.01	-1.24 ± 0.04	-1.79 ± 0.05	9.14 ± 0.00
E10 F9 E9	47625	63686	2.68 0.01	-.93 0.05	-1.76 0.05	7.49 0.00
F9 E9 F8	48122	62938	2.65 0.01	-.79 0.05	-1.86 0.05	8.61 0.00
E9 F8 E8	47580	62261	2.34 0.01	-.88 0.05	-1.46 0.05	4.15 0.00
F8 E8 F7	48073	61509	2.23 0.01	-.82 0.05	-1.41 0.05	7.59 0.01
E8 F7 E7	47536	60825	2.14 0.01	-.48 0.05	-1.66 0.05	9.01 0.01
F7 E7 F6	48023	60069	2.15 0.01	-.89 0.05	-1.26 0.05	9.31 0.01
E7 F6 E6	47491	59392	1.92 0.01	-.43 0.05	-1.49 0.05	10.27 0.01
F6 E6 F5	47973	58641	1.93 0.02	-.73 0.05	-1.20 0.05	10.71 0.01
E6 F5 E5	47447	57953	1.98 0.02	-.64 0.05	-1.35 0.05	12.32 0.01
F5 E5 F4	47923	57204	1.98 0.02	-.63 0.05	-1.35 0.05	12.73 0.01
E5 F4 E4	47402	56511	1.70 0.01	-.69 0.05	-1.00 0.05	8.07 0.01
F4 E4 F3	47873	55774	1.67 0.02	-1.32 0.04	-.35 0.05	10.92 0.01
E4 F3 E3	47357	55070	1.95 0.02	-1.09 0.04	-.86 0.05	15.88 0.01
F3 E3 F2	47823	54336	1.97 0.02	-1.75 0.04	-.22 0.05	15.74 0.00
E3 F2 E2	47311	53604	2.82 0.02	-1.60 0.04	-1.22 0.05	22.57 0.00
F2 E2 F1	47775	52876	2.86 0.02	-1.63 0.04	-1.23 0.05	20.22 0.00
E2 F1 E1	47270	52147	3.14 0.02	-1.74 0.04	-1.40 0.05	21.21 0.00
F1 E1 A1	47888	51200	3.11 0.02	-2.04 0.03	-1.08 0.04	23.86 0.00
E1 A1 A2	47545	50477	2.10 0.03	-1.89 0.05	-.21 0.06	21.73 0.01
A1 A2 D1	48002	49827	2.48 0.05	-2.34 0.09	-.15 0.10	26.13 0.01
A2 D1 C1	47378	49438	2.85 0.03	-2.08 0.05	-.77 0.06	29.87 0.01
D1 C1 D2	47672	48845	2.83 0.03	-2.18 0.05	-.64 0.05	27.11 0.00
C1 D2 C2	47247	48146	2.53 0.03	-2.37 0.04	-.16 0.05	24.38 0.00
D2 C2 D3	47664	47487	2.54 0.03	-1.91 0.05	-.63 0.05	23.38 0.00
C2 D3 C3	47279	46725	2.73 0.03	-1.86 0.04	-.87 0.05	24.80 0.00
D3 C3 D4	47656	46112	2.72 0.03	-1.88 0.05	-.84 0.06	25.01 0.00
C3 D4 C4	47314	45320	2.48 0.03	-1.77 0.05	-.71 0.06	24.18 0.01
D4 C4 D5	47649	44748	2.34 0.04	-1.88 0.05	-.46 0.06	26.84 0.01
C4 D5 C5	47349	43913	2.19 0.04	-1.76 0.05	-.42 0.06	26.40 0.01
D5 C5 D6	47642	43382	2.24 0.04	-1.89 0.05	-.35 0.07	26.22 0.01
C5 D6 C6	47384	42498	2.46 0.04	-1.62 0.05	-.84 0.06	29.69 0.01
D6 C6 D7	47635	42007	2.54 0.05	-1.62 0.06	-.92 0.08	28.84 0.01
C6 D7 C7	47419	41104	2.74 0.05	-2.00 0.06	-.74 0.07	29.69 0.01
D7 C7 D8	47622	40515	2.77 0.05	-2.00 0.06	-.78 0.08	29.38 0.01
C7 D8 C8	47447	39591	1.74 0.05	-1.50 0.06	-.25 0.08	34.62 0.01
D8 C8 D9	47609	39025	1.78 0.06	-1.75 0.07	-.03 0.10	35.44 0.01
C8 D9 C9	47482	38200	1.96 0.07	-2.01 0.07	.05 0.10	30.45 0.01
D9 C9 D10	47604	37654	1.77 0.09	-2.65 0.10	.88 0.14	35.70 0.01
C9 D10 C10	47517	36808	1.52 0.10	-2.26 0.10	.74 0.14	42.80 0.01

\*Azimuth of X-axis is 289.7 degrees. Geographic coordinates of station A1 are: S84°10'52", W154°10'08".

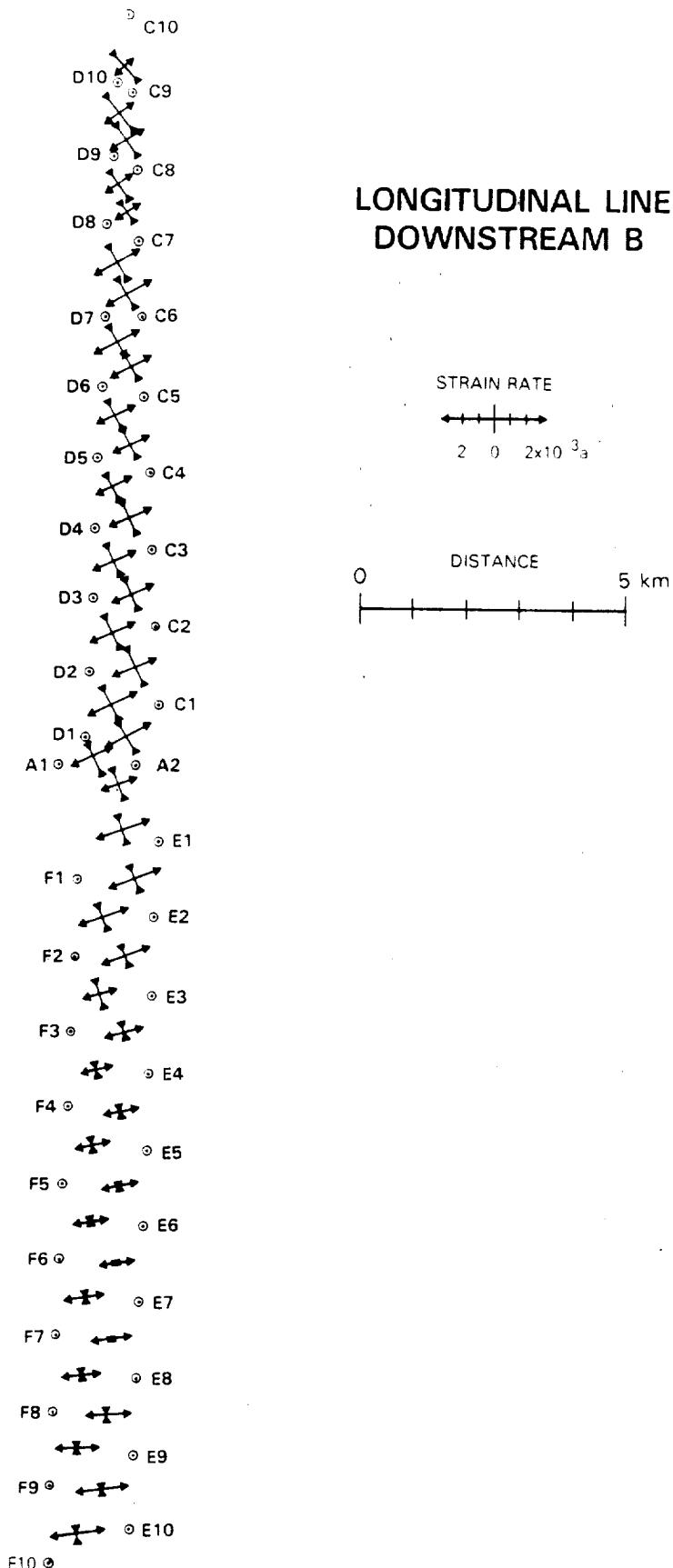


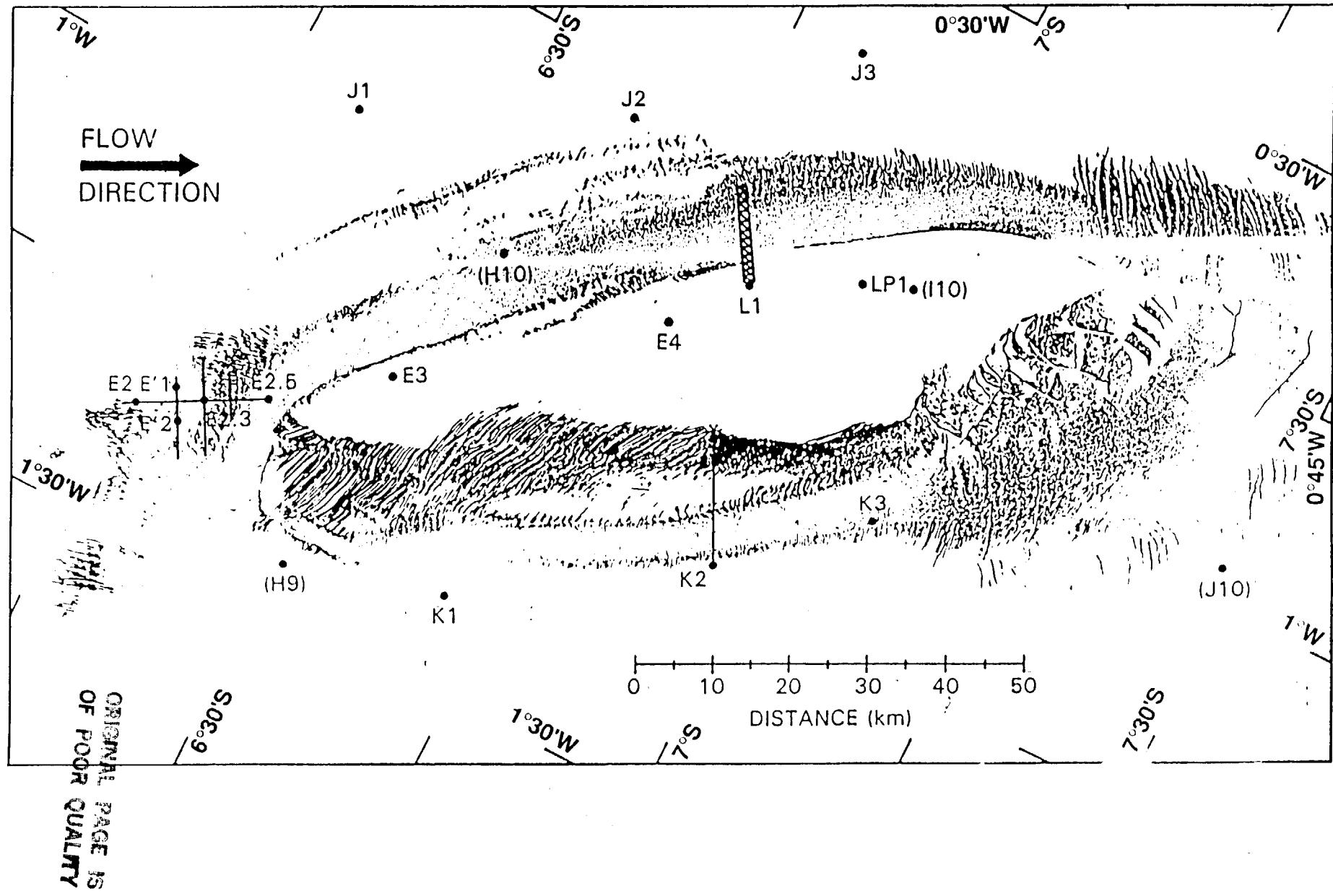
Figure 10: Longitudinal profile of strain rates measured along a double stake line near DNB.

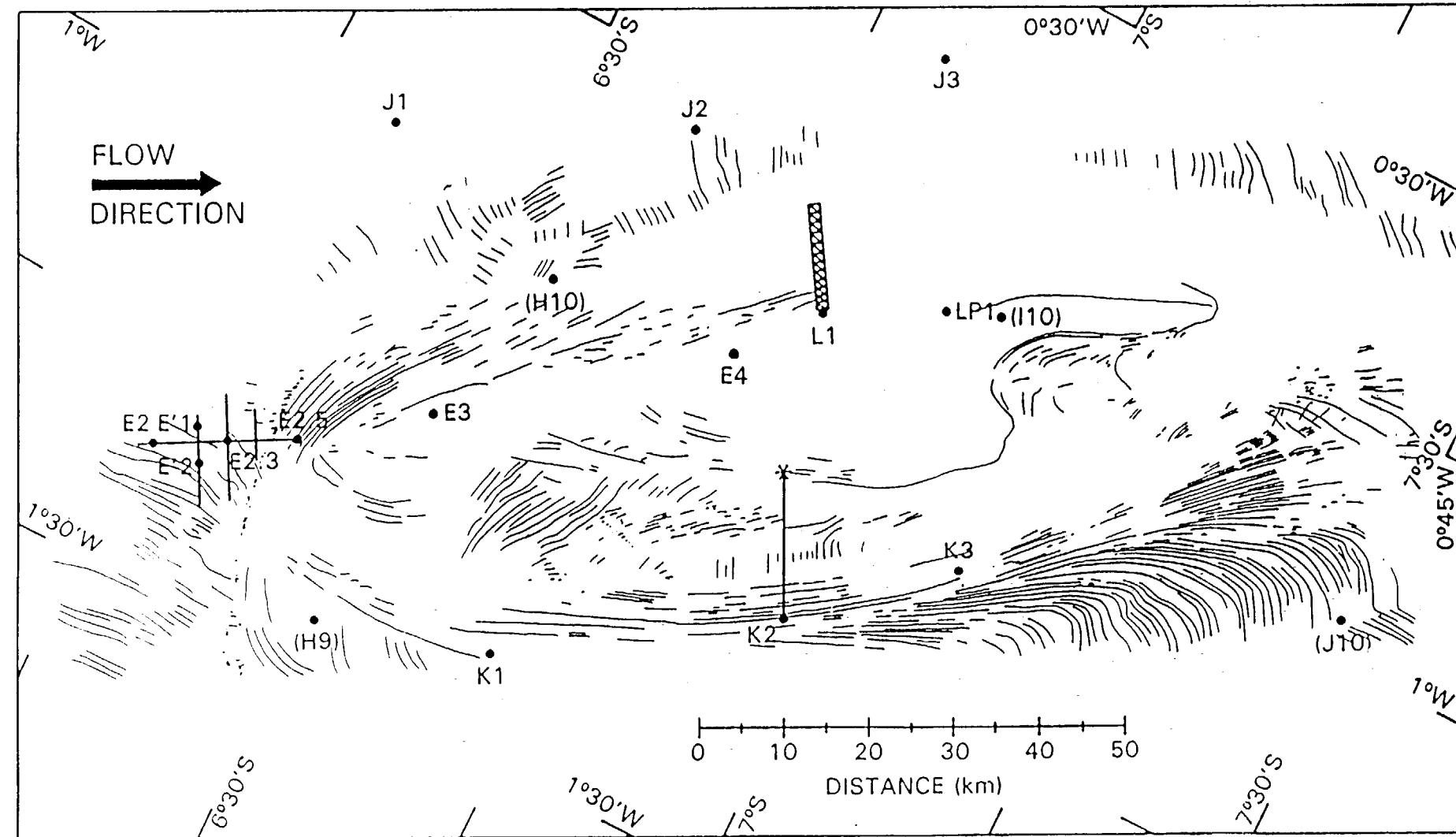
## SECTION 8: AERIAL PHOTOGRAPHY

In January 1985 the U. S. Geological Survey took aerial photographs of the Crary Ice Rise. The photographs covered an area of approximately 175 km by 70 km and were taken at an altitude of 25,000 feet. In all, 570 photographs were obtained. Using every other photograph, a mosaic was created by matching features (stratsugi and crevasses) on adjacent photographs. Overlays were then prepared of crevasses, crests of undulations and open rifts in the ice shelf downstream of the ice rise. Identification of ground control stations provided an estimation of the scale to be 1:47,400. The mosaic and overlays were then photographically reduced to scales of 1:250,000 and 1:500,000. Figures 11(a) and 11(b) show the crevasse and undulation overlays with a geographic grid.

There is a gap in the photographic coverage along a thin strip to the grid northeast of the ice rise. Bad weather forced the data to be collected in two missions separated by 3 days. While an extra photographic strip was taken to provide ample overlay, poor navigation caused errors in position large enough to create the gap. Fortunately, there were two control points on the smaller mosaic which allowed a separate calculation of scale, and allowed this piece to be aligned relative to the larger section. The gap is relatively small and its effect on the mapped features is small. The most apparent difference in the two data sets are the crevasses downstream of the ice rise. In the smaller mosaic, more crevasses are apparent compared to the larger mosaic. This is most likely a result of differing solar azimuth and elevation at the time of each photographic mission.

Figure 11 (a-b): Crevasse and undulation traces from aerial photography of Cray Ice Rise.





## SECTION 9: MULTI-LEG ROSETTES

The multi-leg strain rosette technique outlined in MacAyeal (1985) was used at three stations near the grounding line of the ice stream, N1, N3 and N4. A fourth multi-leg rosette was planted at N2, but it was not resurveyed by the end of the 1985-1986 field season. This technique was chosen because it provides a means of reducing measurement uncertainty under circumstances where the time period between initial station deployment and resurvey is short. Our goal was to recover the strain rate data after approximately 15 days. A redundancy of 40 measurement legs was used by planting 10 outlying stakes and surveying them from four independent central stakes. The details of this innovative method are discussed in Appendix 3. The data are presented in Table 18. Figure 12 shows the similarity of velocities of the ten outlying stakes relative to the velocity of four center stakes. The computed strain rate is indicated in the center of this figure. Figure 13 indicates the scatter of calculated strain rates.

TABLE 18: MULTI-LEG ROSETTE RESULTS

18(a): Comparison between weighted, unweighted and linear method\*

	Single Value Decomposition											
	Weighted Method			Unweighted Method			Linear Method					
	P1	P2	PZ	P1-Azimuth	P1	P2	PZ	P1-Azi	P1	P2	PZ	P1-Azi
				d m d				d m				d m
C3	0.22+0.03	-0.09+0.02	-0.13+0.04	077-33+4.2	0.17	-0.06	-0.17	080-52	0.20	-0.06	-0.14	079
E'1	13.60 0.03	-13.87 0.03	0.27 0.04	070-42 .05	13.59	13.94	0.35	070-36	13.6	-13.9	0.3	070-36
E'2	16.95 0.03	-13.61 0.03	-3.61 0.04	061-06 .02	16.60	-13.77	-2.83	060-04	16.6	-13.8	-2.8	060-06
E2.3	29.89 0.03	-24.43 0.03	-5.46 0.05	076-25 .01	29.41	-22.18	-7.23	074-08	28.8	-22.8	-6.0	074-13
E2.5	28.08 0.03	-37.33 0.03	9.25 0.04	056-31 .01	30.76	-37.99	7.23	060-27	30.9	-37.9	7.0	060-42
E3	1.52 0.03	-0.20 0.03	1.33 0.04	084-56 .78	1.42	-0.20	-1.21	090-54	1.41	-0.24	-1.17	089-24
E4	0.80 0.02	0.14 0.03	-0.95 0.04	342-31 1.8	0.80	0.27	-1.07	344-46	0.80	0.25	-1.05	344
J1	2.81 0.03	-2.45 0.03	-0.36 0.04	069-11 .24	2.86	-2.40	-0.46	068-36	2.88	-2.35	-0.51	068-30
J2	2.79 0.03	-2.85 0.03	0.06 0.04	096-08 .21	2.83	-2.62	-0.21	094-51	2.83	-2.62	-0.21	094-54
J3	1.67 0.03	0.45 0.03	-2.12 0.05	318-16 1.1	1.67	0.44	-2.11	318-42	1.67	0.46	-2.13	318

18(b): Comparison of 3-leg, 10-leg and 40-leg rosette data from N3<sup>+</sup>

3-leg rosettes

	P2	P1	PZ	Pl-Azimuth
				d m
A1 E1 H1	-2.08	2.00	0.75	337-34
B1 F1 I1	-1.45	2.04	-0.60	338-13
C1 G1 J1	-1.64	1.66	-0.02	335-32
A2 E2 H2	-2.18	1.97	0.20	337-10
B2 F2 I2	-1.56	2.12	-0.56	342-11
C2 G2 J2	-1.52	1.52	0.00	334-49
A3 E3 H3	-2.17	1.91	0.26	336-46
B3 F3 I3	-1.54	1.91	-0.37	340-27
C3 G3 J3	-1.59	1.50	0.09	334-20
A4 E4 H4	-2.14	1.89	0.25	336-40
B4 F4 I4	-1.72	2.08	-0.36	240-59
C4 G4 J4	-1.90	2.21	-0.31	335-24
Sample Mean	$-1.8 \pm 0.3$	$1.9 \pm 0.3$	$-0.1 \pm 0.3$	337-30

10-leg rosette

	P2	P1	PZ	Pl-Azimuth
				d m
A1 - J1	-1.69 0.04	1.65 0.03	0.05 0.05	335-51
A2 - J2	-1.68 0.04	1.75 0.03	-0.07 0.05	337-45
A3 - J3	-1.71 0.04	1.65 0.03	-0.06 0.05	334-38
A4 - J4	-1.84 0.04	1.54 0.03	0.30 0.05	335-47
Sample Mean	$-1.73 \pm 0.07$	$1.65 \pm 0.09$	$0.07 \pm 0.16$	336-00

40-leg rosette

	P2	P1	PZ	Pl-Azimuth
				d m
A1 - J40	$-1.73 \pm 0.02$	$1.62 \pm 0.02$	$0.11 \pm 0.03$	335-46

<sup>+</sup>Strain rates are in units of ( $\times 10^{-10} \text{ s}^{-1}$ )

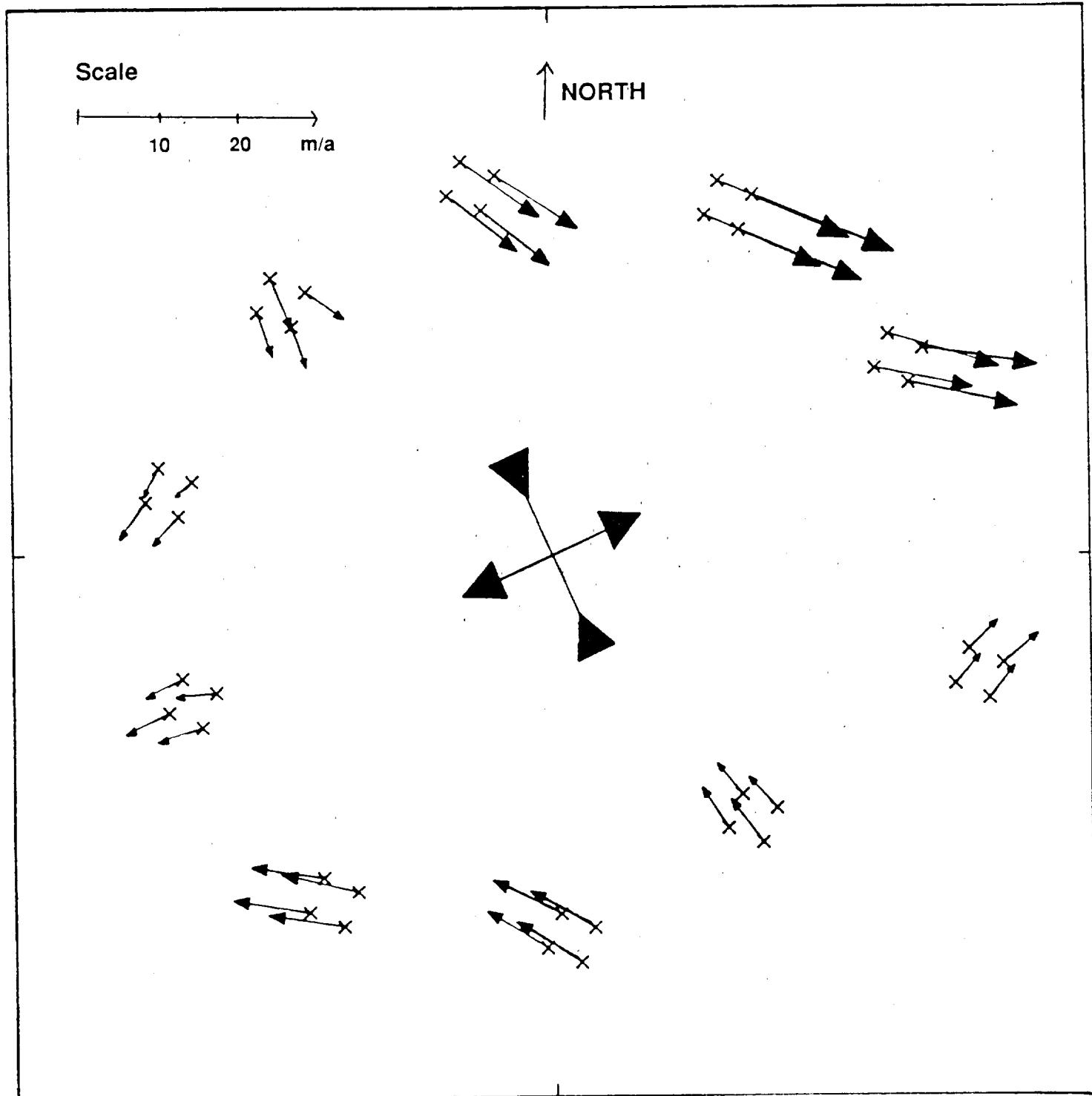
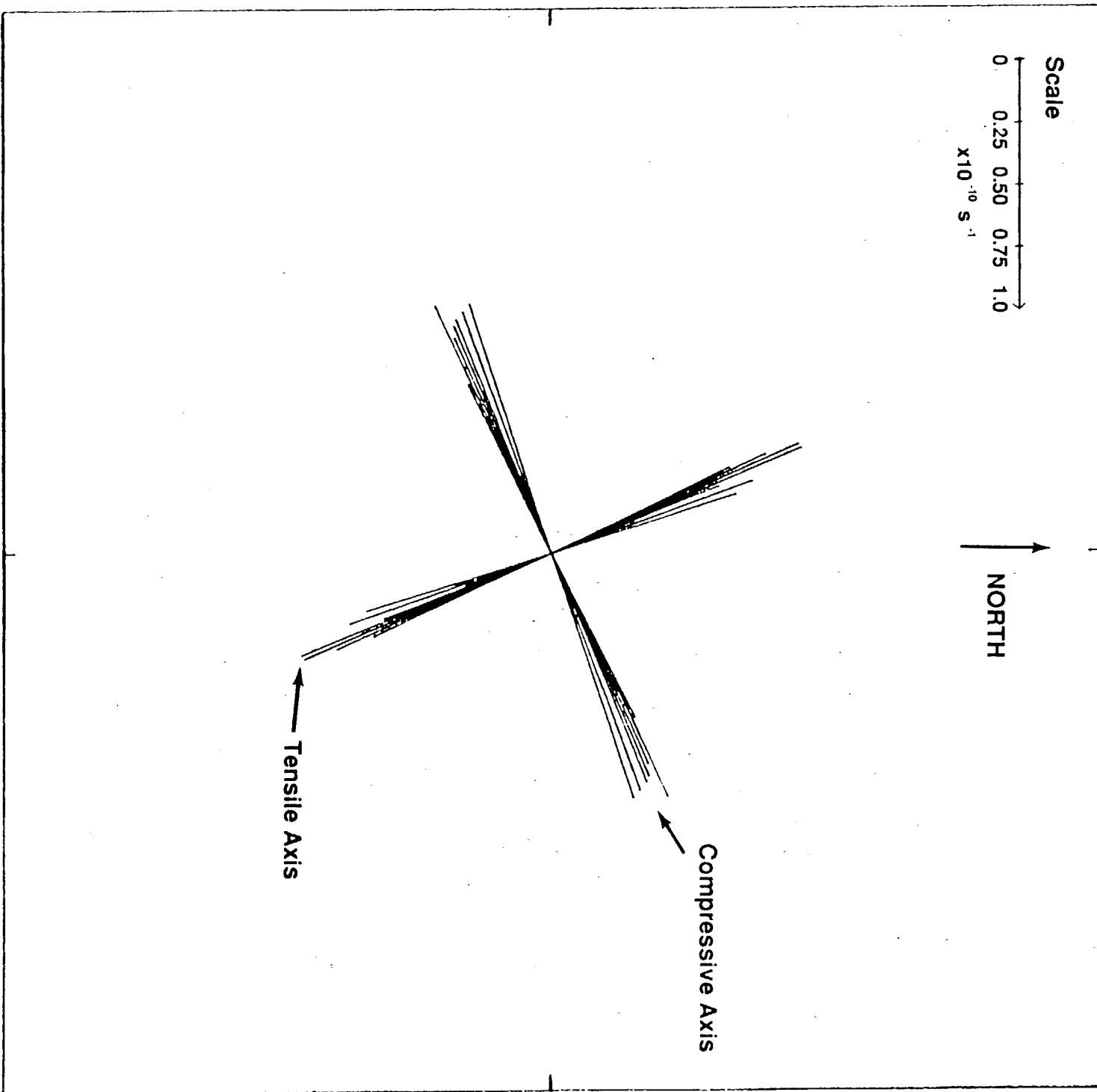


Figure 12-13: Stake movements and strain rates at station N3 from multi-leg rosette results.



#### **ACKNOWLEDGEMENTS**

These data were collected due to the considerable efforts of the following field assistants: James Foster, Richard Otto, John Scofield, Matthew Sturm, David Thompson and Jay Zwally.

Publications from these data (Appendix 4) will never adequately reflect their sacrifices made to this field program.

This research was supported by NSF grants DPP-8207320, DPP-8405287 and DPP-8514543.

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APPENDIX 1 STATION COORDINATES

APPENDIX 2 STATIONS OCCUPIED ONCE WITH GEOCEIVERS BY NASA FIELD PARTIES (AS OF JAN 1986)

APPENDIX 3 DETAILS OF MULTI-LEG ROSETTE TECHNIQUE

APPENDIX 4 SCIENTIFIC PUBLICATIONS

ORIGINAL PAGE IS  
OF POOR QUALITY

APPENDIX 1: STATION COORDINATES

STATION	GEOGRAPHIC			POLAR			STEROGRAPHIC			GRID			SOURCE	
	d	m	s	d	m	s	(km)	(km)	d	m	s	d	m	
CRARY														
CAMP	83 37 15	166 44 30		162.484	-689.595		6 12 33	1 27 47						g2
	83 37 08	166 45 20		162.366	-689.845		6 12 41	1 27 43						g3
O	83 47 14	166 01 28		166.631	-669.539		6 01 44	1 30 02						g2
	83 47 07	166 02 43		166.440	-669.810		6 01 53	1 29 55						g3
A1	83 52 41	165 07 48		174.469	-657.088		5 55 01	1 34 16						s1
A2	83 57 29	164 16 24		181.861	-645.838		5 48 57	1 38 16						g3
B1	83 55 13	165 58 09		163.688	-655.015		5 53 54	1 28 26						s1
	83 55 16	166 00 02		163.307	-655.015		5 53 54	1 28 14						s2
B2	84 02 58	165 56 21		160.540	-640.990		5 46 20	1 26 45						s1
	84 03 02	165 57 31		160.292	-640.924		5 46 18	1 26 36						s2
C1	83 53 00	166 46 17		155.442	-661.244		5 57 16	1 23 59						s2
C2	83 57 32	167 41 16		143.054	-655.433		5 54 08	1 17 18						g3
C3	84 26 08	166 57 08		139.486	-601.891		5 25 15	1 15 22						g3
C4	84 57 52	165 38 44		138.596	-541.585		4 52 42	1 14 54						g2
	84 57 50	165 38 32		138.643	-541.637		4 52 44	1 14 56						g3
D1	83 47 37	167 12 13		152.660	-672.133		6 03 08	1 22 29						s1
	83 47 40	167 13 29		152.392	-672.099		6 03 07	1 22 20						s2
D2	83 47 44	168 22 02		138.936	-674.883		6 04 37	1 15 04						s1
E1	83 42 08	166 52 00		158.921	-681.128		6 07 59	1 25 52						s1
	83 42 25	166 52 29		158.706	-680.638		6 07 43	1 25 44						s2
E2	83 36 38	167 41 37		151.249	-693.317		6 14 34	1 21 43						s1
E1'	83 35 14	168 07 42		146.518	-696.986		6 16 32	1 19 09						e1
E2'	83 33 32	167 52 18		150.301	-699.406		6 17 50	1 21 12						e1
E2.3A	83 33 09	168 12 43		146.290	-700.982		6 18 42	1 19 02						g2
	83 33 05	168 13 14		146.211	-701.128		6 18 46	1 18 59						g3
E2.5	83 30 04	168 39 35		141.932	-707.712		6 22 19	1 16 40						e1
E3	83 22 44	169 34 34		133.057	-723.272		6 30 43	1 11 53						g2
	83 22 45	169 34 30		133.065	-723.238		6 30 42	1 11 53						g3
E4	83 09 20	171 36 38		110.924	-752.137		6 46 16	0 59 55						g2
	83 09 20	171 36 40		110.917	-752.138		6 46 16	0 59 55						g3
F1	83 39 43	166 01 13		170.048	-683.055		6 09 01	1 31 52						s1
	83 39 44	165 58 18		170.620	-682.881		6 08 56	1 32 11						s2
F2	83 31 59	166 02 37		173.229	-697.045		6 16 34	1 33 35						s1
	83 31 55	166 02 57		173.192	-697.182		6 16 38	1 33 34						s2
G1	83 41 38	165 06 18		180.024	-676.816		6 05 39	1 37 16						s1
G2	83 35 38	164 23 28		191.437	-685.239		6 10 12	1 43 25						g3
G3	83 25 42	162 43 09		216.822	-696.957		6 16 30	1 57 08						g2
	83 25 33	162 43 45		216.783	-697.260		6 16 40	1 57 06						g3
G4	83 16 40	161 26 55		237.555	-707.867		6 22 23	2 08 19						g2
	83 16 32	161 27 17		237.558	-708.127		6 22 31	2 08 20						g3
H1	83 47 02	164 48 42		180.862	-666.220		5 59 56	1 37 43						s1
	83 46 58	164 48 44		180.888	-666.342		6 00 00	1 37 44						s2
H2	83 46 09	163 40 13		194.558	-664.058		5 58 46	1 45 07						g3
J1	83 35 26	171 35 33		104.081	-704.199		6 20 26	0 56 14						g2
	83 35 19	171 37 00		103.816	-704.457		6 20 34	0 56 05						g3
J2	83 19 04	173 04 26		89.503	-736.795		6 38 00	0 48 21						g2
	83 18 56	173 05 32		89.297	-737.069		6 38 09	0 48 14						g3
J3	83 07 22	174 55 29		67.580	-760.927		6 51 01	0 36 30						g2
	83 07 12	174 56 21		67.415	-761.252		6 51 11	0 36 25						g3
K1	83 10 24	168 09 09		155.683	-742.140		6 40 52	1 24 06						g2

	83 10 18	168 09 44	155.595	-742.348	6 40 59	1 24 03	g3
K2	82 56 35	169 57 59	136.582	-771.946	6 56 56	1 13 46	g2
	82 56 29	169 58 35	136.479	-772.152	6 57 03	1 13 43	g3
K3	82 49 26	171 09 53	122.447	-787.748	7 05 28	1 06 08	g2
	82 49 21	171 10 36	122.306	-787.926	7 05 33	1 06 03	g3
L1	83 06 12	172 25 21	101.022	-759.397	6 50 11	0 54 34	g2
	83 06 11	172 25 23	101.018	-759.429	6 50 12	0 54 34	g3
LP1	83 00 04	172 59 57	94.761	-771.674	6 56 48	0 51 11	g2
	83 00 05	172 59 53	94.772	-771.641	6 56 47	0 51 11	g3

#### DOWNSTREAM B

DNB	84 10 28	154 18 43	280.431	-582.938	5 14 59	2 31 31	g2
	84 10 28	154 21 24	279.950	-583.169	5 15 06	2 31 16	g3
A11	84 18 16	154 30 54	272.116	-570.887	5 08 29	2 27 02	e2
A19	84 24 27	154 42 21	265.316	-561.428	5 03 23	2 23 21	g2
	84 24 26	154 45 03	264.888	-561.664	5 03 30	2 23 08	g3
B10	84 02 42	154 07 06	288.659	-594.954	5 21 28	2 35 58	e2
B18	83 57 00	153 33 32	299.160	-601.571	5 25 02	2 41 38	g2
	83 56 58	153 35 42	298.809	-601.815	5 25 10	2 41 27	g3
B25	83 52 02	153 10 38	307.331	-607.811	5 28 23	2 46 02	g2
	83 52 02	153 10 42	307.305	-607.790	5 28 23	2 46 02	g3
C10	84 12 53	153 06 35	290.548	-572.943	5 09 35	2 36 60	e2
E10	84 10 44	155 37 12	266.821	-588.749	5 18 07	2 24 10	e2
E19	84 09 03	156 51 07	255.328	-597.219	5 22 42	2 17 58	g2
	84 09 01	156 53 31	254.935	-597.454	5 22 49	2 17 45	g3
G1	84 03 58	152 08 59	307.847	-582.632	5 14 48	2 46 19	g2
	84 03 58	152 11 46	307.358	-582.860	5 14 56	2 46 04	g3
G2	84 00 28	150 32 03	327.325	-579.350	5 13 02	2 56 51	g2
	84 00 31	150 34 49	326.813	-579.533	5 13 08	2 56 35	g3
H2	83 53 49	150 25 25	334.528	-589.442	5 18 28	3 00 44	g2
	83 53 50	150 25 29	334.502	-589.422	5 18 28	3 00 44	g3
M0	84 17 46	158 10 58	235.383	-587.986	5 17 43	2 07 11	g2
	84 17 42	158 13 27	235.004	-588.271	5 17 52	2 06 59	g3
M1	84 22 05	156 53 27	245.440	-575.170	5 10 48	2 12 38	e3
M2	84 25 40	158 07 53	230.456	-574.185	5 10 16	2 04 32	g3
M3	84 21 12	159 26 11	220.227	-587.036	5 17 13	1 59 00	g3
M4	84 13 28	159 29 32	224.677	-600.678	5 24 34	2 01 24	s2
M5	84 10 01	158 15 49	239.874	-601.663	5 25 06	2 09 37	s2
M6	84 13 17	157 05 06	249.841	-591.024	5 19 21	2 14 60	s2
N1	83 50 30	161 56 00	212.095	-650.186	5 51 17	1 54 35	g3
N2	83 36 11	160 22 29	238.619	-669.187	6 01 31	2 08 55	g3
N3	83 46 55	157 07 35	268.417	-636.248	5 43 45	2 25 01	g3
N4	84 03 44	159 01 17	236.066	-615.660	5 32 39	2 07 33	g3

#### DOWNSTREAM C

DNC	82 49 18	152 26 55	368.858	-707.021	6 21 51	3 19 13	g2
A1	82 59 48	152 28 50	359.458	-689.941	6 12 39	3 14 09	g2
A2	83 10 43	152 30 45	349.723	-672.169	6 03 05	3 08 54	g2
A3	83 18 42	152 32 05	342.620	-659.161	5 56 04	3 05 05	g2
B1	82 38 48	152 25 46	378.117	-724.178	6 31 06	3 24 12	g2
B2	82 30 50	152 24 33	385.225	-737.155	6 38 05	3 28 02	g2
B3	82 18 07	152 23 07	396.480	-757.922	6 49 16	3 34 06	g2
C1	82 49 11	151 21 23	382.372	-700.051	6 18 06	3 26 31	g2
C2	82 49 02	150 15 28	395.862	-692.833	6 14 12	3 33 48	s1
C3	82 42 46	149 06 00	415.755	-694.676	6 15 10	3 44 32	g2

D19	82	48	55	155	17	34	333.619	-725.099	6	31	37	3	00	11	g2
H5	82	35	24	153	15	02	370.542	-735.160	6	37	01	3	20	07	g2

WISCONSIN STATIONS ON ROSS ICE SHELF

C2	83	57	32	167	41	16	143.054	-655.433	5	54	08	1	17	18	g3
W3	84	55	53	155	44	05	231.253	-512.998	4	37	15	2	04	59	
W4	84	49	29	161	02	39	186.643	-543.411	4	53	41	1	40	52	g3
W4525	84	56	05	154	37	46	240.947	-508.107	4	34	36	2	10	13	g3
W5	84	35	31	166	37	35	138.884	-584.169	5	15	41	1	15	03	g3
W6	84	25	14	169	21	55	114.329	-608.869	5	29	01	1	01	47	g3
W9	84	19	10	159	46	37	218.037	-591.872	5	19	49	1	57	49	g3

SOURCES:

g2: GEOCEIVER FIX FROM 1984-1985 FIELD SEASON  
 g3: GEOCEIVER FIX FROM 1985-1986 FIELD SEASON  
 s1: SAT NAV FIX FROM 1983-1984 FIELD SEASON  
 s2: SAT NAV FIX FROM 1984-1985 FIELD SEASON  
 e1: EXTRAPOLATION USING GEOC. POSITION OF E2.3A AND SKIDOO ODOMETER  
 e2: EXTRAPOLATION USING GEOC. POSITION OF DNB AND SKIDOO ODOMETER  
 e3: EXTRAPOLATION USING GEOC. POSITION OF MO

APPENDIX 2: STATIONS OCCUPIED ONCE WITH GEOCEIVERS BY NASA  
FIELD PARTIES (AS OF JANUARY 1986)

SITE	LATITUDE	LONGITUDE	EL (m)	PA	TIME	DATE
CRARY						
C3	84 26	07.625	166 57	08.402	56 47	15 35
H2	83 46	09.281	163 40	14.034	50 25	20 00
A2	83 57	28.588	164 16	24.332	48 58	3 35
G2	83 35	37.797	164 23	28.932	39 12	13 8
F1	83 39	27.587	166 02	54.77	40 24	7 45
DOWNSTREAM	B					
M2	84 25	40.249	158 07	52.608	70 34	06 00
M3	84 21	12.548	159 26	10.863	56 8	29 10
N1	83 50	29.663	161 56	00.344	49 27	11 40
N2	83 36	11.064	160 22	29.464	52 28	00 45
N3	83 46	55.426	157 07	35.203	55 50	20 00
N4	84 03	44.258	159 01	16.731	70 6	05 59
DOWNSTREAM	C					
O	82 49	17.859	152 26	55.258	99 25	21 51
A1	82 59	48.191	152 28	49.899	91 24	1 38
A2	83 10	42.763	152 30	44.793	86 6	7 40
A3	83 18	42.206	152 32	04.592	74 17	19 25
B2	82 30	49.841	152 24	33.203	63 16	23 44
B3	82 18	07.056	152 23	07.012	107 1	3 3
C1	82 49	10.783	151 21	22.814	116 16	20 52
C3	82 42	45.673	149 06	00.324	124 8	7 36
F1	82 48	54.844	155 17	33.593	30 42	20 13
WISCONSIN STATIONS ON ROSS ICE SHELF						
W6	84 25	13.521	169 21	54.822	50 23	18 28
W5	84 35	30.675	166 37	34.879	59 5	05 24
W4525	84 56	04.558	154 37	46.514	94 20	00 56
C2	83 57	31.509	167 41	15.866	45 7	02 05
W9	84 19	10.379	159 46	36.782	62 22	23 16
W4	84 49	28.752	161 02	39.241	72 56	10 4

### APPENDIX 3: DETAILS OF MULTI-LEG ROSETTE TECHNIQUE

Four multi-leg rosettes were planted between 1st December and 6th December, 1985 by two independent two-man survey teams. At each site, four central stakes (labeled 01 to 04) were planted into the snow using post levels to maintain verticality. These four central stakes were positioned to form the corners of a square with 150-m sides. Ten outlying stakes were positioned in an array surrounding the central square. Their distance from the central square was approximately 1500 m, and they were separated from each other by an angle of 36 degrees about 01.

Distances were measured between each of the four central stakes and the ten outlying stakes using a Geodimeter 112 EDM with a nominal accuracy of 0.01 m over 1500 m; the times of each measurement were recorded to the nearest minute. Rounds of angles between a reference outlying stake, 'A' and the other 9 outlying stakes were measured using a Wild T-2 theodolite to a nominal accuracy of 3 seconds of arc, mounted on a tripod optically plumbed over the center of each of the four central stakes. The commencement time of each round was recorded. The true azimuth of the 'A' stake from each of the four central stakes was measured using observations onto the sun. The Magnavox MX1502 geoceiver provided both the time and the longitude used for the calculation of azimuth. The estimated error is 15 seconds of arc. This is greater than the error in observing angles between the outlying stakes, but the azimuth is only used to determine the relative vorticity (rigid body rotation). The geoceiver was also used to determine the geodetic position of one of the central stakes to an estimated error of 20 meters, (see Section 1).

Resurvey of three of the four multi-leg rosettes was accomplished using the same measurement techniques by one two-man survey party between 16 December and 19 December, 1985. After the resurvey, three outlying stakes and one central stake were extended so that the station could be resurveyed in the future as a 3-leg rosette.

## Data Organization

The observed relative position of the  $N$ th outlying stake with respect to each of the four central stakes is described in polar coordinates ( $R_n, \theta_n, t_{R_n}, t_{\theta_n}$ ) where  $R_n$  denotes the radial distance,  $\theta_n$  denotes the true azimuth (determined by adding the azimuth of the 'A' stake to the angular separation between the 'A' stake and the  $N$ th stake), and  $t_{R_n}$  and  $t_{\theta_n}$  denote the times at which  $R_n$  and  $\theta_n$  are observed.

The four sets of ten relative positions were combined to comprise one set of 40 relative positions. The effect of lumping the data together in this manner is to average out any strain rate gradients within the 150-m box formed by the four central stakes. This averaging is not considered detrimental because our primary objective is to determine the average strain rate within the whole rosette and not the strain rate gradient.

Another modification made to the data was that the initial observations of the 'A' stake's azimuth were substituted for the final observations of the azimuth made during the resurvey. This was done to avoid introducing the azimuth uncertainty into the determination of stake displacements. Effectively, the azimuth data is used to determine the average positions of the stakes with respect to each other, but not to determine their relative displacements. Comparisons were made between strain rate reductions in which this substitution was and was not done, and no significant differences (they were less than 0.01 times the standard deviation of the strain rate components) were noticed. The relative vorticity determined as a result of this substitution must be corrected by adding the relative vorticity associated with the 'rigid rotation' determined by the rate of change of the true azimuth. We describe the uncertainty associated with the determination of the relative vorticity separately, below.

Once two sets of 40 stake positions representing the initial and final survey are established, average positions  $\bar{R}_n$ ,  $\bar{\Theta}_n$  and relative velocities  $\dot{R}_n$ ,  $\dot{\Theta}_n$  are computed:

$$\bar{R}_n = \frac{1}{2}(R_n^i + R_n^f) \quad (1)$$

$$\bar{\Theta}_n = \frac{1}{2}(\Theta_n^i + \Theta_n^f) \quad (2)$$

$$\dot{R}_n = \frac{R_n^f - R_n^i}{t_{R_n}^f - t_{R_n}^i} \quad (3)$$

$$\dot{\Theta}_n = \frac{\Theta_n^f - \Theta_n^i}{t_{\Theta_n}^f - t_{\Theta_n}^i} \quad (4)$$

where superscripts i and f denote values determined during the initial and final surveys, respectively.

These data are next converted to average positions  $\bar{x}_n$ ,  $\bar{y}_n$  and relative velocities  $\dot{x}_n$ ,  $\dot{y}_n$  in a coordinate system having an X-axis aligned with true North and a Y-axis aligned with true West,

$$\bar{x}_n = \bar{R}_n \cos \bar{\Theta}_n \quad (5)$$

$$\bar{y}_n = -\bar{R}_n \sin \bar{\Theta}_n \quad (6)$$

$$\dot{x}_n = \dot{R}_n \cos \bar{\Theta}_n - \bar{R}_n \sin \bar{\Theta}_n \dot{\Theta}_n \quad (7)$$

$$\dot{y}_n = -\dot{R}_n \sin \bar{\Theta}_n - \bar{R}_n \cos \bar{\Theta}_n \dot{\Theta}_n \quad (8)$$

### Determination of Strain Rates

The determination of strain rates from a 40-leg, over-determined stake velocity data set requires a least-squares procedure known as singular value decomposition. Here we outline the details of this calculation.

The strain rates are determined from the gradients in  $u$  and  $v$  where  $u$  and  $v$  are components of the velocity in the  $x$  and  $y$  directions, respectively (following Jaeger and Cook, 1976). The first step is to determine these velocity gradients by solving for  $S$ , the least-squares estimate of the velocity gradient column vector, in the following equation. (Note that here we write the equations in terms of a 40-leg rosette; the equations are easily modified to address an  $N$ -leg rosette where  $N$  is arbitrary.)

$$\underline{\underline{A}} \times \underline{S} = \underline{d} - \underline{z} \quad (9)$$

where,

$$\underline{A} = \begin{bmatrix} \bar{x}_1 & \bar{y}_1 & 0 & 0 \\ \vdots & \vdots & \vdots & \vdots \\ \bar{x}_{40} & \bar{y}_{40} & 0 & 0 \\ 0 & 0 & \bar{x}_1 & \bar{y}_1 \\ \vdots & \vdots & \vdots & \vdots \\ 0 & 0 & \bar{x}_{40} & \bar{y}_{40} \end{bmatrix} \quad \underline{S} = \begin{bmatrix} \frac{\partial u}{\partial x} \\ \frac{\partial u}{\partial y} \\ \frac{\partial v}{\partial x} \\ \frac{\partial v}{\partial y} \end{bmatrix}$$

$$\underline{d} = \begin{bmatrix} \dot{x}_1 \\ \vdots \\ \dot{x}_{40} \\ \dot{y}_1 \\ \vdots \\ \dot{y}_{40} \end{bmatrix} \quad \underline{z} = \begin{bmatrix} z_1 \\ \vdots \\ z_{40} \end{bmatrix}$$

and  $z$  represents the measurement error in the observed values of  $d$ . We assume that the covariance of  $\underline{z}$  is given by:

$$i = 1, 40 \quad \text{cov}(z_i, z_j) = \left\{ \frac{\sigma_R^2 \cos^2 \bar{\Theta}_i}{(t_{R_i}^f - t_{R_i}^l)^2} + \frac{\sigma_{\Theta_i}^2 \bar{R}_i^2 \sin^2 \bar{\Theta}_i}{(t_{\Theta_i}^f - t_{\Theta_i}^l)^2} \right\} \hat{\sigma}_{ij} \quad (10)$$

$$i = 41,80 \quad \text{cov}(z_i, z_j) = \left\{ \frac{\sigma_R^2 \sin^2 \bar{\Theta}_i}{(t_{R_i}^f - t_{R_i}^i)^2} + \frac{\sigma_{\Theta}^2 \bar{R}_i^2 \cos^2 \Theta_i}{(t_{\Theta_i}^f - t_{\Theta_i}^i)^2} \right\} \delta_{ij}$$

where  $\delta_{ij}$  is the Kroneker delta,  $\sigma_R$  is the standard deviation estimate of the measurements of  $R_n$  in meters, and  $\sigma_{\Theta}$  is the standard deviation estimate of the measurement of  $\Theta_n$  in radians. This assumes covariance is consistent with the assumption that each of the 40 measurements of the stake velocities are statistically independent. However, the measurements are not completely independent because certain types of error associated with disturbing outlying stakes can affect four of the data points, one each associated with the observation from each central stake. We assume that these statistical dependencies produce only small off-diagonal terms in the covariance matrix of  $z$  compared to the diagonal terms. As stated previously, we assume that  $\sigma_R = 0.01$  m and  $\sigma_{\Theta} = 3$  sec in radians.

To correct for some components of  $z$  being larger than others by virtue of the inequality of  $\sigma_R$  and  $\bar{R}_n \sigma_{\Theta}$ , we weight the data matrix  $\underline{d}$  and the position matrix  $\underline{\underline{A}}$ ,

$$d_i^w = d_i \left[ \frac{\sigma^2}{\text{cov}(z_i, z_i)} \right]^{1/2} \quad (11)$$

$$A_{ij}^w = A_{ij} \left[ \frac{\sigma^2}{\text{cov}(z_i, z_j)} \right]^{1/2} \quad (12)$$

and,

$$z_i^w = z_i \left[ \frac{\sigma^2}{\text{cov}(z_i, z_i)} \right]^{1/2} \quad (13)$$

where the superscript w's denote weighted elements, and

$$\sigma^2 = \frac{1}{80} \sum_{i=1}^{80} \text{cov}(z_i, z_i) \quad (14)$$

Equation (9) now becomes,

$$\underline{A}^w \times \underline{S} = \underline{d}^w + \underline{z}^w \quad (15)$$

where the covariance matrix of  $\underline{z}^w$  is now  $\sigma^2$  on the diagonal.

To invert  $\underline{A}^w$  for determining  $S_{ij}$ , the four eigenvalues  $\lambda_k$  and the eigenvectors  $\underline{r}_k$  of  $[\underline{A}^w]^T [\underline{A}^w]$  are computed:

$$\lambda_1 = \frac{\alpha_1 + \gamma_1}{2} + \left\{ \frac{(\alpha_1 + \gamma_1)^2}{4} - (\alpha_1 \gamma_1 - \beta_1^2) \right\}^{1/2} \quad (16)$$

$$\lambda_2 = \frac{\alpha_1 + \gamma_1}{2} - \left\{ \frac{\alpha_1 + \gamma_1}{4} - (\alpha_1 \gamma_1 - \beta_1^2) \right\}^{1/2} \quad (17)$$

$$\lambda_3 = \frac{\alpha_2 + \gamma_2}{2} - \left\{ \frac{(\alpha_2 + \gamma_2)^2}{4} - (\alpha_2 \gamma_2 - \beta_2^2) \right\}^{1/2} \quad (18)$$

$$\lambda_4 = \frac{\alpha_2 + \gamma_2}{2} - \left\{ \frac{(\alpha_2 + \gamma_2)^2}{4} - (\alpha_2 \gamma_2 - \beta_2^2) \right\}^{1/2} \quad (19)$$

and,

$$\underline{r}_1^T = \left( \frac{-\beta_1}{\alpha_1 - \lambda_1} \left( 1 + \frac{\beta_1^2}{(\alpha_1 - \lambda_1)^2} \right)^{-1/2}, \left( 1 + \frac{\beta_1^2}{(\alpha_1 - \lambda_1)^2} \right)^{-1/2}, 0, 0 \right) \quad (20)$$

$$\underline{r}_2^T = \left( \frac{\alpha_1 - \lambda_1}{\beta_1} \left( 1 + \frac{(\alpha_1 - \lambda_1)^2}{\beta_1^2} \right)^{-1/2}, \left( 1 + \frac{(\alpha_1 - \lambda_1)^2}{\beta_1^2} \right)^{-1/2}, 0, 0 \right) \quad (21)$$

$$\underline{r}_3^T = \left( 0, 0, -\frac{\beta_2}{\alpha_2 - \lambda_3} \left( 1 + \frac{\beta_2^2}{(\alpha_2 - \lambda_3)^2} \right)^{-1/2}, \left( 1 + \frac{\beta_2^2}{(\alpha_2 - \lambda_3)^2} \right)^{-1/2} \right) \quad (22)$$

$$\underline{r}_4^T = \left( 0, 0, \frac{\alpha_2 - \lambda_3}{\beta_2} \left( 1 + \frac{(\alpha_2 - \lambda_3)^2}{\beta_2^2} \right)^{-1/2}, \left( 1 + \frac{\beta_2^2}{(\alpha_2 - \lambda_3)^2} \right)^{-1/2} \right) \quad (23)$$

and where,

$$[\underline{\underline{A}}^w]^T [\underline{\underline{A}}^w] = \begin{bmatrix} \alpha_1 & \beta_1 & 0 & 0 \\ \beta_1 & \gamma_1 & 0 & 0 \\ 0 & 0 & \alpha_2 & \beta_2 \\ 0 & 0 & \beta_2 & \gamma_2 \end{bmatrix} \quad (24)$$

with

$$\alpha_1 = \sum_{i=1}^{40} \left[ \frac{\sigma^2}{\text{cov}(z_i, z_i)} \right] \bar{x}_i^2 \quad (25)$$

$$\beta_1 = \sum_{i=1}^{40} \left[ \frac{\sigma^2}{\text{cov}(z_i, z_i)} \right] \bar{x}_i \bar{y}_i \quad (26)$$

$$\gamma_1 = \sum_{i=1}^{40} \left[ \frac{\sigma^2}{\text{cov}(z_i, z_i)} \right] \bar{y}_i^2 \quad (27)$$

$$\alpha_2 = \sum_{i=1}^{40} \left[ \frac{\sigma^2}{\text{cov}(z_{i+40}, z_{i+40})} \right] \bar{x}_i^2 \quad (28)$$

$$\beta_2 = \sum_{i=1}^{40} \left[ \frac{\sigma^2}{\text{cov}(z_{i+40}, z_{i+40})} \right] \bar{x}_i \bar{y}_i \quad (29)$$

and,

$$\gamma_2 = \sum_{i=1}^{40} \left[ \frac{\sigma^2}{\text{cov}(z_{i+40}, z_{i+40})} \right] \bar{y}_i^2 \quad (30)$$

The least-squares estimate of  $\underline{S}$  is found by constructing the inverse of  $\underline{A}$ ,

$$[\underline{A}^w]^{-1} = \underline{R} \cdot \underline{\Gamma}^{-1} \cdot \underline{Q}^T \quad (31)$$

where  $R$ ,  $\Gamma$  and  $Q$  are matrices defined in MacAyeal (1985). In component notation,

$$\begin{aligned} S_i = & \sum_{k=1}^4 r_i^k \frac{1}{\lambda_k} \left\{ \sum_{i=1}^{40} \left( \bar{x}_i r_1^k + \bar{y}_i r_2^k \right) \frac{\sigma^2}{\text{cov}(z_i, z_i)} d_i \right. \\ & \left. + \sum_{i=41}^{80} (\bar{x}_{i-40} r_3^k + \bar{y}_{i-40} r_4^k) \frac{\sigma^2}{\text{cov}(z_i, z_i)} d_i \right\} \end{aligned} \quad (32)$$

where the superscript  $k$  refers to the eigenvector number.

The principal axes  $e_1$  and  $e_2$  of the horizontal strain rate tensor are determined from  $\underline{S}$  by using the relations (Jaeger and Cook, 1976)

$$\Theta = \frac{1}{2} \tan^{-1} \left( \frac{S_2 + S_3}{S_1 + S_4} \right) \quad (33)$$

$$\dot{e}_1 = S_1 \cos^2 \Theta + \frac{1}{2} (S_2 + S_3) \sin 2\Theta + S_4 \sin^2 \Theta \quad (34)$$

and

$$\dot{e}_2 = S_1 \sin^2 \Theta - \frac{1}{2} (S_2 + S_3) \sin 2\Theta + S_4 \cos^2 \Theta \quad (35)$$

where  $\Theta$  is the counterclockwise angle (in radians) between the x-axis, or north, and the  $e_1$ -axis.

#### Computation of Error

Perhaps the most critical element of our analysis is the computation of expected uncertainty for the derived values of  $\Theta$ ,  $\dot{e}_1$ ,  $\dot{e}_2$ , and  $\dot{e}_{zz} = -\dot{e}_1 - \dot{e}_2$ . This is critical for two reasons. Firstly, the short time period over which the 40-leg rosette was allowed to deform means that the observed stake displacements may be close to the limits of detectability, and secondly, the value of one of the strain rate principal components may be several orders of magnitude less than the maximum component and its value may be statistically insignificant.

The expected uncertainty, or estimated standard deviation of the components of  $\underline{S}$ , are determined formally through the singular value decomposition procedure (MacAyeal, 1985).

$$\sigma_{S_i} = \left[ \text{cov}(S_i - S_i^*, S_i - S_i^*) \right]^{1/2} \quad (36)$$

where  $S^*$  is composed of the 'true' value of the velocity gradients, and  $\underline{S}$  is, as stated before, the least-squares estimate of  $S^*$ . Observe that the components of  $\sigma$  depend only on the covariance of  $z^w$ , the eigenvectors and the eigenvalues of  $[A^w]^T [A^w]$ ; the latter two of which depend only on stake positioning. In component notation,

$$\sigma_{S_i} = \left\{ \sum_{k=1}^4 \sum_{i=1}^4 \frac{r_i^T r_k^T}{\lambda_k} \cdot \sigma^2 \right\}^{1/2} \quad (37)$$

where  $\sigma^2$  is given by Equation (14), and subscripts  $i$  refer to eigenvector number.

The uncertainty of  $\Theta$ ,  $\dot{e}_1$ ,  $\dot{e}_2$  are computed from the  $\sigma_{S_i}$ 's using the relationship for functions of random variables (Boas, 1983).

$$\sigma_\Theta = \frac{1}{2} \left( 1 + \frac{S_2 - S_3}{S_1 + S_4} \right) \left( \frac{S_2 + S_3}{(S_1 - S_4)^2} \right)^2 (\sigma_{S_1}^2 + \sigma_{S_4}^2) + \left( \frac{1}{S_1 - S_4} \right)^2 (\sigma_{S_2}^2 + \sigma_{S_3}^2)^{1/2} \quad (38)$$

$$\sigma_{\dot{e}_1} = \left\{ \cos^4 \Theta \sigma_{S_1}^2 + \frac{1}{4} \sin^2 2\Theta (\sigma_{S_2}^2 + \sigma_{S_3}^2) + \sin^4 \Theta \sigma_{S_4}^2 + ((S_4 - S_1) \sin 2\Theta + (S_2 + S_3 \cos 2\Theta)^2) \sigma^2 \right\}^{1/2} \quad (39)$$

and,

$$\sigma_{e_2} = \left\{ \sin^4 \Theta \sigma_{S_1}^2 + \frac{1}{4} \sin^2 2\Theta (\sigma_{S_2}^2 + \sigma_{S_3}^2) + \cos^4 \Theta \sigma_{S_4}^2 + ((S_1 - S_4) \sin 2\Theta - (S_2 + S_3) \cos 2\Theta)^2 \sigma_{\text{int}}^2 \right\}^{1/2} \quad (40)$$

### **Data reduction programs**

The above technique for reducing 40-leg rosette data can be used to reduce any rosette design (including the standard 3-leg rosette), or other strain figures which yield data in the form  $\bar{x}_i$ ,  $\bar{y}_i$ ,  $x_i$ , and  $y_i$ . A reliable calculator program which will reduce rosettes having up to 100 legs (a limit imposed by the size of calculator memory) has been developed and tested. This program was tested by two means. First, the results of reducing several 3-leg rosettes were compared with results using other methods, and second, multi-leg rosette data were synthesized using a known strain rate, and then reduced by the program to test whether the known strain rate is reproduced. This program, available on request, requires the following equipment: Hewlett-Packard 41-CX hand calculator having "Date" and "Time" functions and at least 3 modules for memory extension, an HP magnetic card reader to input the program, and an HP thermal printer to verify correct data input. The use of a hand calculator was chosen to allow data reduction in the field and this battery powered equipment operated well within the tented shelters used during the field program.

### **Results**

Here, we present information comparing the SVD method with other methods of data reduction. First, the SVD method is compared to the method used in Section 2 of this report for calculating

strain rates from 3-leg rosette data to check the reliability of both these methods. Second, subsets of the 40-leg rosette data are used, which illustrate that the derived strain rate converges as the number of legs is increased.

Table 18 presents strain rates derived by both methods from 3-leg strain rosettes surrounding the Crary Ice Rise. To illustrate the effect of weighting, Table 18 also includes results obtained when the data is not weighted in the SVD method. It is seen that data weighting may be considered unnecessary in the analysis of 3-leg rosette data because errors in the distance and angle measurements are insignificant compared to actual changes in those quantities over the survey period. For all 14 stations, the principal horizontal strain rate components derived by the two methods differed by less than 10 percent and fall within the computed uncertainty limits obtained by both methods. This comparison serves as a useful check on both our data-analysis techniques.

To illustrate convergence of the SVD technique in producing an accurate estimate of the horizontal strain rates from the 40-leg rosette N3, twelve independent 3-leg rosettes and four independent 10-leg rosettes from the 40-leg rosette data are synthesized by considering subsets of the stake array. The results are given in Table 18. There is considerable scatter in the results of the twelve 3-leg rosette as expected from the short time period before which the 3-leg rosettes were resurveyed, but the strain rate falls within the sample standard deviation of the strain rate derived from the analysis of the 40-leg data. The sample standard deviation of the twelve 3-leg rosettes results is larger than the confidence limits computed from 40-leg rosette data. This could have resulted from three factors: 1) our estimates of measurement error on  $R$  and  $\theta$  are too low, 2) the sample of twelve 3-leg rosettes is too small to produce an accurate sample standard deviation, and 3) there is strain rate variation within the confines of our stake array that is averaged out by the 40-leg analysis. All these possible

explanations have some validity. However, the possible underestimate of the measurement accuracy seriously affects the scientific conclusions of our study; certainly any underestimate is not a factor of 10, which is needed if that alone causes the discrepancy.

As another illustration of the multi-leg rosette analysis, four 10-leg rosette sub-sets were analyzed, each of the 4 central stakes being the center of a separate stake array. There is much less scatter in these results (Table 18) when compared to the twelve 3-leg rosettes. The sample standard deviation is also consistent with the confidence limits of the 40-leg rosette showing there is a convergence in both the derived strain rate and the confidence limits as the number of stakes is increased.

#### INSTRUMENT CORRECTIONS

Benchmark comparisons were made periodically between all EDM's used in the 1985-1986 field season. These comparisons revealed that 1 EDM used in the initial and final surveys of 40-leg stations N1 and N4 requires a  $-5.5 \times 10^{-5} + 0.1 \times 10^{-5}$  parts-per-part correction to the initial measured distance.

#### MULTI-LEG ROSETTE UTILITY

Multi-leg rosettes can serve a useful purpose under circumstances when more practical 3-leg rosettes will not provide sufficient accuracy over the time interval between survey and resurvey. Consideration must be taken of the large commitment in time required to deploy these rosettes. Experienced surveyors, in good weather, needed 1.5 days to deploy a 40-leg rosette and 1 day to resurvey it. The method also requires instrument dependability.

#### APPENDIX 4: SCIENTIFIC PUBLICATIONS

- 1) Bindschadler, R.A., B. Koci, S. Shabtaie, and E.P. Roberts, in press. Evolution of Crary Ice Rise, Antarctica, Annals of Glaciology, 12.
- 2) Bindschadler, R.A., P.L. Vornberger, S.N. Stephenson, E.P. Roberts, S. Shabtaie, and D.R. MacAyeal, in press. Ice-Shelf Flow at the Boundary of Crary Ice Rise, Annals of Glaciology, 11.
- 3) Bindschadler, R.A., D.R. MacAyeal, and S.N. Stephenson, 1987. Ice Stream-Ice Shelf Interaction in West Antarctica. In The Dynamics of the West Antarctic Ice Sheet (C.J. Van der Veen and J. Oerlemans, eds.), D. Reidel Pub. Co., p. 161-180.
- 4) Bindschadler, R.A., S.N. Stephenson, D.R. MacAyeal, and S. Shabtaie, 1987. Ice Dynamics at the Mouth of Ice Stream B, Antarctica, J. Geophys. Res., 92, No. B9, p. 8885-8894.
- 5) Lindstrom, D.R., submitted. West Antarctic Ice Sheet Formation, Annals of Glaciology, 11.
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- 8) MacAyeal, D.R., R.A. Bindschadler, K.C. Jezek, and S.

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- 11) Stephenson, S.N. and R.A. Bindschadler, 1988. Observed Velocity Fluctuations on a Major Antarctic Ice Stream, Nature 334.
- 12) Thomas, R.H., S.N. Stephenson, R.A. Bindschadler, S. Shabtaie, and C.R. Bentley, in press. Thinning and Grounding Line Retreat on the Ross Ice Shelf. Annals of Glaciology, 11.



## Report Documentation Page

1. Report No	2. Government Accession No.	3. Recipient's Catalog No
NASA TM-100708		
4. Title and Subtitle	5. Report Date	
Data Report for the Siple Coast Project	October 1988	
7. Author(s)	6. Performing Organization Code	
R. A. Bindschadler, S. N. Stephenson, E. P. Roberts, D. R. MacAyeal, and D. R. Lindstrom	671.0	
9. Performing Organization Name and Address	8. Performing Organization Report No	
Goddard Space Flight Center Greenbelt, Maryland 20771	88B0214	
12. Sponsoring Agency Name and Address	10. Work Unit No.	
National Aeronautics and Space Administration Washington, D.C. 20546-0001	11. Contract or Grant No.	
15. Supplementary Notes	13. Type of Report and Period Covered	
S. N. Stephenson - Science Applications Research, Lanham, Maryland, 20706; E. P. Roberts - Department of Geology, University of Maryland, College Park, Maryland, 20742; D. R. MacAyeal and D. R. Lindstrom - Department of Geophysical Sciences, University of Chicago, Chicago, Illinois, 60637.	Technical Memorandum	
16. Abstract	14. Sponsoring Agency Code	
This report presents data collected during three field seasons of glaciological studies in Antarctica and describes the methods employed. The region investigated covers the mouths of Ice Streams B and C (the Siple Coast) and Cray Ice Rise on the Ross Ice Shelf. Measurements included in the report are: surface velocity and deformation from repeated satellite geociever positions; surface topography from optical levelling; radar sounding of ice thickness; accumulation rates; near-surface densities and temperature profiles, and mapping from aerial photography.		
17. Key Words (Suggested by Author(s))	18. Distribution Statement	
Ice Sheet Glaciology Ice Dynamics	Unclassified - Unlimited	
Subject Category 43		
19. Security Classif. (of this report)	20. Security Classif. (of this page)	21. No. of pages
Unclassified	Unclassified	22. Price





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